

Application of artificial intelligence to estimate the reference evapotranspiration in sub-humid Doon valley

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Abstract Estimation of evapotranspiration (ET) is an essential component of the hydrologic cycle, which is also requisite for efficient irrigation water management planning and hydro-meteorological studies at both the basin and catchment scales. There are about twenty well-established methods available for ET estimation which depends upon various meteorological parameters and assumptions. Most of these methods are physically based and need a variety of input data. The FAO-56 Penman–Monteith method (PM) for estimating reference evapotranspiration (ET_0) is recommend for irrigation scheduling worldwide, because PM generally yields the best results under various climatic conditions. This study investigates the abilities of artificial neural networks (ANN) to improve the accuracy of monthly evaporation estimation in sub-humid climatic region of Dehradun. In the first part of the study, different ANN models, comprising various combinations of training function and number of neurons were developed to estimate the ET_0 and it has been compared with the Penman–Monteith (PM) ET_0 as the ideal (observed) ET_0 . Various statistical approaches were considered to estimate the model performance, i.e. Coefficient of Correlation (r), Sum of Squared Errors, Root Mean Square Error, Nash–Sutcliffe Efficiency Index (NSE) and Mean Absolute Error. The ANN model with Levenberg–Marquardt training algorithm, single hidden layer and nine number of neuron schema was found the best predicting capabilities for the study station with Coefficient of Correlation (r) and NSE

value of 0.996 and 0.991 for calibration period and 0.990 and 0.980 for validation period, respectively. In the subsequent part of the study, the trend analysis of ET_0 time series revealed a rising trend in the month of March, and a falling trend during June to November, except August, with more than 90% significance level and the annual declining rate was found to 1.49 mm per year.

Keywords Evapotranspiration · Penman–Monteith · Neural networks · Mann–Kendall test

Introduction

Evapotranspiration (ET) is one of the major components to identify the behaviour of hydrologic cycle. Correct estimation of reference evapotranspiration (ET_0) is necessary for determining irrigation demands, irrigation scheduling, water resources management, and environmental impact assessment, water balance studies at regional and local levels, and modelling of various rainfall-runoff and ecosystem models. Further, the actual ET can be determined using crop coefficients, which is mainly a function of various crop characteristics and local environmental conditions. This paper deals only with the estimation of ET_0 . Inaccurate estimates of ET_0 can lead to unproductive use of water, false calibrations of hydrological models, and erratic assessments of groundwater recharge. There are many empirical equations available for estimating ET_0 , but their cogency is sometimes restricted to particular geographic locations or/and specific climatic conditions. Allen et al. (1998) defined ET_0 as “the rate of evapotranspiration from a hypothetical crop with an assumed crop height (0.12 m) and a fixed canopy resistance (70 s/m) and albedo

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(0.23) which would closely resemble evapotranspiration from an extensive surface of green grass cover of uniform height, actively growing, completely shading the ground and not short of water. The Food and Agriculture Organization (FAO) of the United Nations recently adopted a standardized form of the Penman–Monteith equation (FAO56 PM) in an effort to provide a common, globally valid standard for estimating ET_0 , developing crop coefficients, and evaluation/calibration of other ET_0 methods when Lysimeter measurements are unavailable (Allen et al. 1998). There are more than twenty established methods available for estimating of ET_0 , these can be broadly classified into three groups temperature-based approaches (i.e. Hargreaves–Samani and Thornthwaite), radiation based approaches (i.e. Priestley–Taylor and Turc) and combination approach (i.e. Penman–Monteith). Most of these methods are physically based and need a variety of input data.

Apart from the tradition modeling techniques, artificial intelligence techniques such as artificial neural networks (ANNs), have proven their feasibility in modeling complex nonlinear phenomena in the recent past (Kumar et al. 2002; Kisi 2006a, b; Rahimikhoob 2010; Kumar et al. 2011; Baba et al. 2013; Goyal et al. 2014; Shamshirband et al. 2015). ANN application to ET modeling initiated when Odhiambo et al. (2001) validated the accurate estimate daily ET_0 with the fuzzy ET_0 models when compared with FAO56 PM. Further, Kumar et al. (2002) reported the precise ET_0 estimations using the developed ANN models with various ANN architectures.

It has been demonstrated that neural networks are a competitive substitute to conventional models for modeling nonlinear systems. Recently, artificial neural networks (ANN) have been applied in meteorological and agro-ecological modelling; most of the applications reported in the literature concern estimation, prediction and classification problems (Dogan 2008; Kim and Kim 2008; Parasuraman et al. 2007; Zanetti et al. 2007; Keskin and Terzi 2006; Sudheer et al. 2003; Trajkovic et al. 2003). Neural network applications have gained popularity due to their functional characteristics, lesser data requirement and capability of long term forecasting which provide many advantages over traditional analytical approaches. There has been a rising drift of application of ANNs in water resources and hydrologic modeling. The basic goal of this paper is to evaluate the comparative performance of ANN based approach with the Penman–Monteith method (Allen et al. 1998) for estimating the ET_0 on monthly basis. The trend analysis along with the magnitude of trend of ET_0 for the period of 34 years also has been done for the study station located at Doon valley of Shivaliks hills.

Materials and methods

Study area and data used

Monthly climatic data used for ET_0 estimation were acquired from meteorological observatory from Forest Research Institute (FRI), Dehradun. The observatory is located in the Doon Valley of Shivaliks range of the Himalayas. Its geographical location is 30.20 N latitude and 77.59 E longitude. Dehradun has an altitude of 629 m above mean sea level. It has sub-humid, sub-tropic climate. The summer is too humid and hot, the winter is too cold and the rainy season has a heavy rainfall. The monthly mean of maximum temperature lies in the range of 18.6–38.6 °C. The minimum temperature varies between 1.7 and 23.4 °C. June is the hottest and January is the coolest month. The monsoon season experiences about 87% of the average annual rainfall of about 2041.9 mm. Location map of study area is shown in Fig. 1.

While modeling, the monthly temperature (Min., Max., and Avg.), relative humidity (Min. and Max.), average wind speed, sunshine hours and rainfall data from 1980 to 2013 (i.e. 408 datasets) were divided into calibration period 1980–2004 (300 datasets) and validation period 2005–2013 (108 datasets).

Penman–Monteith method

Penman–Monteith (PM) FAO56 method for estimating ET_0 is a physically based method that unambiguously integrates both physiological and aerodynamic parameters (Xu et al. 2006). It is supposed to be the most consistent and worldwide accepted method for ET_0 estimation under various types of climate. The most common form of the PM method in computing ET_0 is given as (Allen et al. 1998):

$$ET_0 = \frac{0.408\Delta(R_n - G) + \gamma \frac{900}{T+273} u_2 (e_s - e_a)}{\Delta + \gamma(1 + 0.34u_2)} \quad (1)$$

where ET_0 is the reference evapotranspiration (mm/day); R_n is the net radiation at the crop surface ($MJ/m^2/d$); G is the soil heat flux density ($MJ/m^2/d$); T is the mean daily air temperature (°C); u_2 is the wind speed at a 2 m height above the ground (m/s); e_s is the saturation vapour pressure (kPa); e_a is the actual vapour pressure (kPa); $e_s - e_a$ is the saturation VPD (kPa); Δ is the slope of vapour pressure versus temperature curve at temperature T (kPa/°C), and γ is the psychrometric constant (kPa/°C). The reference crop was assumed as green grass with an albedo of 0.23.

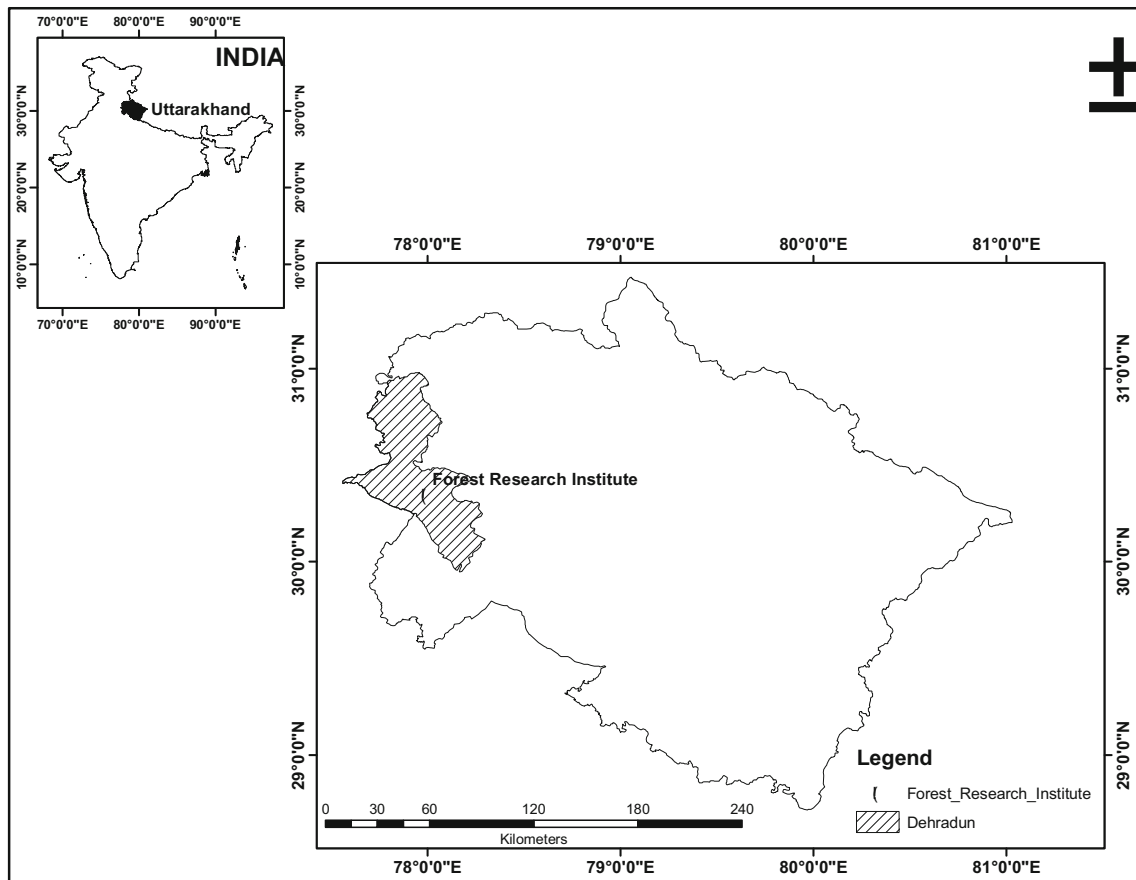


Fig. 1 Location map of study area

Artificial neural networks (ANNs)

Commonly an ANN architecture consist of a set of the artificial neuron inputs or signals (x), a weighted average calculation of them (z) using the summation function and weights (w), and some activation function (f) to produce an output, where

$$Z = \sum_{i=1}^n x_i w_i \tag{2}$$

The connections between the input layer and the middle or hidden layer contain weights [Eq. (2)], which are usually determined through training the system. The hidden layer sums the weighted inputs and uses the transfer function to create an output value. The transfer function is a relationship between the internal activation level of the neuron (called activation function) and the outputs. A typical transfer function is the sigmoid function, which varies from 0 to 1 for a range of inputs (Melesse et al. 2011). The 408 dataset was divided into training and validation data (300 training and 108 validations) for single hidden layer in three-layer feed forward back propagation algorithm. The feed forward back propagation algorithms

was applied using different transfer function and best transfer function (logsig) is chosen on the basis of performance of model. The supervised learning of model is done using reference evapotranspiration (ET_0) as output and temperature (Min., Max., and Avg.), relative humidity (Min. and Max.), average wind speed, and sunshine hours as input parameters.

Levenberg–Marquardt back propagation

The Levenberg–Marquardt also known as the damped least-squares (DLS) method is used to solve non-linear least squares problems. It is a trust region-based method with a hyper-spherical trust region (Ajmera and Goyal 2012; Kumar et al. 2015). The Levenberg–Marquardt algorithm uses this approximation to the Hessian matrix in the following Newton-like update

$$X_{k+1} = X_k - [J_k^T J + \mu I]^{-1} J_k^T e \tag{3}$$

where J = Jacobian matrix, which contains the first derivatives of the network errors regarding the weights and biases; e = vector of network errors; and I = identity matrix.

When the scalar μ is zero, this is just a Newton's method, using the approximate Hessian matrix. When μ is large, this becomes gradient descent with a small step size. Newton's method is faster and more accurate near an error minimum, so the aim is to shift toward Newton's method as quickly as possible (Hagan and Menhaj 1994 and MATLAB online reference manual). Thus, μ is decreased after each successful step (reduction in performance function) and is increased only when a tentative step would increase the performance function.

Goodness of fit of model performance The performance of the any model can be evaluated in terms of variety of statistical criteria. However, four standard statistical evaluation criteria, viz. Coefficient of Correlation (r), Sum of Squared Errors (SSE), Root Mean Square Error (RMSE), Nash–Sutcliffe Efficiency Index (NSE) and Mean Absolute Error (MAE) were used to measure the performance of models (Moriassi et al. 2007).

1. Coefficient of correlation (r)

$$r = \frac{n(\sum x_0 \times x_f) - (\sum x_0) \times (\sum x_f)}{\sqrt{[n \sum x_0^2 - (\sum x_0)^2][n \sum x_f^2 - (\sum x_f)^2]}} \quad (4)$$

where x_0 = observed data, x_f computed data, n = total number of events considered.

2. Sum of Squared Errors (SSE)

$$SSE = \sum_{i=1}^n (x_0 - \bar{x})^2 \quad (5)$$

where x_0 = observed data, x_f computed data.

3. Nash–Sutcliffe Efficiency (NSE):

$$NSE = 1 - \frac{\sum_{i=1}^n (x_0 - x_f)^2}{\sum_{i=1}^n (x_0 - \bar{x}_0)^2} \quad (6)$$

where x_0 = observed data, x_f computed data, \bar{x} = mean of the observed data, and n = total number of events considered.

4. Mean Absolute Error (MAE)

$$MAE = 1 - \frac{\sum_{i=1}^n |x_0 - x_f|}{N} \quad (7)$$

where x_0 = observed data, x_f computed data, \bar{x} = mean of the observed data, and N = number of non-missing data points.

Trend analysis

The trend detection in the ET₀ time series of the FRI, Dehradun station has been done using the most commonly used non-parametric Mann–Kendall (MK) and Modified Mann–Kendall (MMK) test (Mann 1945; Kendall and Stuart 1961; Kendall 1962). It is a rank-based technique, which is good enough to accommodate the effect of intermediate extremes events and take cares of the skewed variables. The rate of the change in the time series at the said station has also been determined with Sen' Slope estimator method (Sen 1968). According to this test, the null hypothesis H₀ states that the de-seasonalized data ($x_1 \dots x_n$) is a sample of n independent and identically distributed random variables. The alternative hypothesis H₁ of a two-sided test is that the distributions of x_k and x_j are not identical for all $k, j \leq n$ with $k \neq j$. The test statistic S , which has mean zero and a variance computed by Eq. (4), is calculated using Eqs. (2) and (3), and is asymptotically normal (Hirsch et al. 1982):

$$S = \sum_{k=1}^{n-1} \sum_{j=k+1}^n \text{sgn}(x_j - x_k) \quad (8)$$

$$\text{Sgn}(x_j - x_k) = \begin{cases} +1 & \text{if } (x_j - x_k) > 0 \\ 0 & \text{if } (x_j - x_k) = 0 \\ -1 & \text{if } (x_j - x_k) < 0 \end{cases} \quad (9)$$

$$\text{Var}(S) = \left[n(n-1)(2n+5) - \sum_t t(t-1)(2t+5) \right] / 18 \quad (10)$$

The notation t is the extent of any given tie and \sum_t denotes the summation over all ties. In cases where the sample size $n > 10$, the standard normal variate z is computed using Eq. (5). In a two-sided test for trend, H₀ should thus be accepted if $|z| \leq z_{\alpha/2}$ at the α level of significance. A positive value of 'S' indicates an 'upward trend'; likewise, a negative value of 'S' indicates 'downward trend':

$$Z_{mk} = \begin{cases} \frac{S-1}{\sqrt{\text{Var}(S)}} & \text{if } S > 0 \\ 0 & \text{if } S = 0 \\ \frac{S+1}{\sqrt{\text{Var}(S)}} & \text{if } S < 0 \end{cases} \quad (11)$$

The Mann–Kendall test has two parameters that are of importance to trend detection. These parameters are the significance level that indicates the trends strength and the slope magnitude estimates that indicates the direction as well as the magnitude of the trend. The MK test checks the null hypothesis of no trend versus the alternative hypothesis of the existence of increasing or decreasing trend. If Z_{mk} is positive, then the trend is increasing, and if Z_{mk} is negative, then the trend is decreasing.

Trend analysis using the MMK test

The modified Mann–Kendall test is similar to MK test, except the effect of all significant autocorrelation coefficients has to be removed from the time series (Hamed and Rao 1998; Basistha et al. 2009) and its empirical significance level are more accurate. For this purpose, a modified variance of S, designated as $\text{Var}(S)^*$, was used as follows:

$$\text{Var}(S)^* = V(S) \frac{n}{n^*} \tag{12}$$

where n^* = effective sample size. The n/n^* ratio was computed directly from the equation proposed by Hamed and Rao (1998) as

$$\frac{n}{n^*} = 1 + \frac{2}{n(n-1)(n-2)} \sum_{i=1}^{n-1} (n-i)(n-i-1)(n-i-2)\gamma_i \tag{13}$$

where n = actual number of observations, and r_i = lag- i significant auto coefficient of correlation of rank i of time series. Once $\text{Var}(S)^*$ was computed from Eq. (6), then it was substituted for $\text{Var}(S)$ in Eq. (5). The estimate Z-values is then tested for significance of trend. The normal threshold levels, for various significance levels are as follows: 1.645 for 10%, 1.96 for 5% and 2.33 for 1% level of significance.

Magnitude of the trend

The magnitude of the trend or the rate of change in a time series was estimated using a non-parametric Sen’s slope estimator method (Sen 1968). In this method, a linear trend in the time series is generally assumed and the slope (T_i) of all pairs are first calculated by

$$T_i = \frac{x_j - x_k}{j - k} \quad \text{for, } i = 1, 2, 3 \dots N \tag{14}$$

where X_j and X_k area data values at time j and k ($j > k$), respectively. The median of these N values of T_i is Sen’s estimator of slope which is calculated as

$$\beta = \begin{cases} T_{\frac{N+1}{2}} & N \text{ is odd} \\ \frac{1}{2} (T_{\frac{N}{2}} + T_{\frac{N+2}{2}}) & N \text{ is even} \end{cases} \tag{15}$$

The trend is supposed to be upward (increasing) with positive values of B and vice versa.

Result and discussion

ANN model for ET₀

Initially the ANN Models have been tested with the different training algorithm and six numbers of neurons and

single hidden layer. To get the more accurate ANN model the performance of six predictive training algorithms was evaluated (Table 1) while developing the model itself. The most promising training algorithm in the hidden layer of ANN model was determined by trial and error, at which a model perform better.

Table 1 indicates that model with training algorithm Levenberg–Marquardt algorithm (trainlm) resulted satisfactory value of coefficient of correlation (r) as 0.997 and 0.989 during training and validation, respectively. The figures in bold represents the best performing indicies. The model performance indices; SSE and RMSE with scaled estimate and target is lowest as 0.733 and 0.049 for training and 1.102 and 0.101 for validation periods, respectively. Predictive power of ANN models, i.e. NSE has been worked out and found highest as 0.994 and 0.978 for training and validation runs, respectively, for model with trainlm function. MAE was found lowest in case tarincrgf function. Based on the four out of five performance indices the trainlm function was chosen for further modeling.

Increasing the number of neurons in the hidden layer, the network gets an over fit, that problem has to be resolved. To get the optimum number of neurons at which network should have to perform its best, trial and error method has been applied. Optimization of the number of neurons for further modeling exercise is one of the most crucial part of any ANN model development. The model with learning function “trainlm” trained with 75% of data has been assessed for optimum number of neurons. Neurons in the intermediate hidden layer have been varied from 1 to 10. The comparison of different performance parameters presented in Table 2 for the neuron variation. It can be identified that model with “trainlm” algorithm and nine neurons performed best for the validation results for most of the statistical indices. Scatterplots between observed (FAO56-PM) and modelled (ANN) ET₀ for validation and calibration runs can be shown in Fig. 2 demonstrating the model performance. It can be observed from these scatterplots that ANN model has a strong predictive power for the given data sets and ET₀ for the FRI, station Dehradun ranges from 2.4 to 5.0 mm/day.

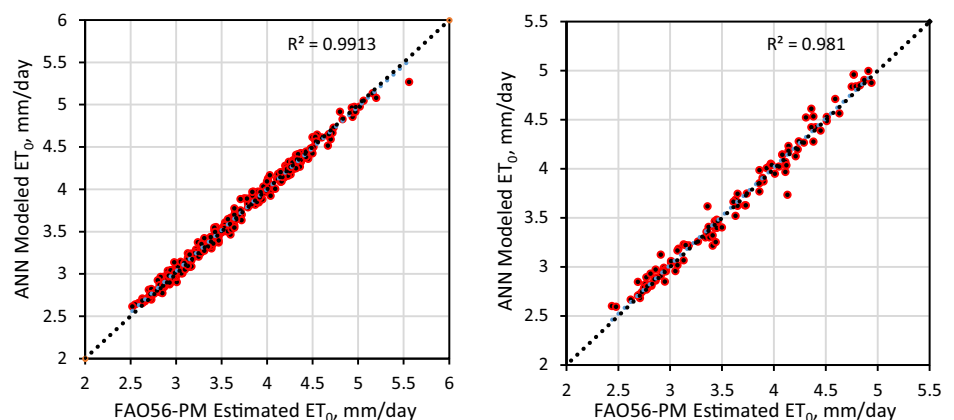
The values of mean monthly FAO-PM ET₀ and predicted ET₀ with the ANN model with “trainlm” training algorithm at nine neurons can be compared as shown in Table 3. It can be perceived that the value of ET₀ for the pre-monsoon months, i.e. summer season is higher due to hot and humid weather conditions and the lowest ET₀ values were appearing during the winter months. The total annual ET₀ was observed to be 1300.85 mm. And it is found comparable and close to the model predicted annual ET₀ value of 1306.55 mm. Figure 3 represents the basic statistic of ET₀ for the study period of 34 years (1980–2013) in the form of box plot.

Table 1 Performance of different training algorithms for ANN based ET_0 modeling

SN	Training functions	Training					Validation				
		r	SSE	RMSE	NSE	MAE	r	SSE	RMSE	NSE	MAE
1	trainbfg	0.997	0.785	0.051	0.994	0.926	0.983	1.760	0.128	0.964	0.873
2	trainbr	0.997	0.851	0.053	0.993	0.922	0.976	2.508	0.152	0.949	0.861
3	traincgb	0.996	1.090	0.060	0.991	0.913	0.988	1.233	0.107	0.975	0.878
4	traincgf	0.992	2.088	0.083	0.983	0.876	0.969	3.047	0.168	0.938	0.843
5	traincgp	0.994	1.413	0.069	0.989	0.903	0.977	2.315	0.146	0.953	0.885
6	trainlm	0.997	0.733	0.049	0.994	0.930	0.989	1.102	0.101	0.978	0.866

Table 2 Performance of ANN model with the varied number of neurons

SN	No. of neurons	Training					Validation				
		r	SSE	RMSE	NSE	MAE	r	SSE	RMSE	NSE	MAE
1	1	0.979	5.161	0.131	0.958	0.814	0.976	2.909	0.164	0.941	0.809
2	2	0.990	2.564	0.092	0.979	0.869	0.968	3.341	0.176	0.932	0.843
3	3	0.994	1.413	0.069	0.989	0.908	0.958	4.391	0.202	0.911	0.838
4	4	0.996	0.976	0.057	0.992	0.917	0.978	2.399	0.149	0.951	0.841
5	5	0.995	1.369	0.068	0.989	0.904	0.941	6.314	0.242	0.872	0.828
6	6	0.995	1.369	0.068	0.989	0.904	0.941	6.314	0.242	0.872	0.828
7	7	0.995	1.369	0.068	0.989	0.904	0.941	6.314	0.242	0.872	0.828
8	8	0.997	0.776	0.051	0.994	0.929	0.965	3.530	0.181	0.929	0.852
9	9	0.996	1.155	0.062	0.991	0.910	0.990	0.978	0.095	0.980	0.882
10	10	0.997	0.848	0.053	0.993	0.926	0.978	2.229	0.144	0.955	0.865

Fig. 2 Performance of ANN model with “trainlm” training algorithm at nine neurons for ET_0 estimation for calibration (left) and validation (right) runs

Trend analysis

The annual, seasonal, and monthly time series of ET_0 for the study station were examined for the possible trend detection. The Z-statistic values of monotonic trend test were obtained through the nonparametric Mann–Kendall and Modified Mann–Kendall tests. As a precautionary step,

the significant lag-1 serial correlation effect has been removed from the time series of ET_0 before the trend testing. The FRI station, Dehradun observed significant annual ET_0 decreases at a rate of 1.49 mm/year on annual basis.

Table 4 represents the trend statistics MK_Z and MMK_Z with the magnitude of the trend (Sen’s slope). It

Table 3 FAO-PM and ANN estimated mean monthly ET₀ (mm)

Month	PM (ET ₀)	ANN
Jan	84.08	85.50
Feb	83.97	85.25
Mar	119.38	119.88
Apr	135.00	134.80
May	142.12	143.68
Jun	119.00	122.46
Jul	91.41	93.74
Aug	96.48	95.26
Sep	113.66	109.84
Oct	124.41	124.90
Nov	101.23	100.55
Dec	90.10	90.68
Annual	1300.85	1306.55

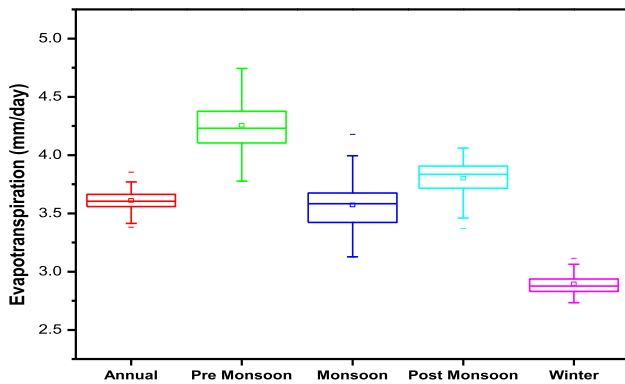


Fig. 3 Box plots for the annual and seasonal ET₀

can be observed from the Table 4 that the FRI station, Dehradun is showing an increasing trend in the month of March, and a decreasing trend during June to November, except August, with more than 90% significance. The monsoon and post-monsoon season are witnessing the falling ET₀ trend at a rate of 1.03 and 0.51 mm/year with a significance level of 95 and 99%, respectively, whereas the winter and pre-monsoon seasons are not having any trends. Overall, the station is supporting a decreasing trend in the ET₀ at 95% significance level with the rate of 1.49 mm/year on annual basis.

Conclusion

Correct estimation of evapotranspiration is prerequisite for the precise planning and management of irrigation schemes, hydrological modeling, groundwater recharge, crop performance, and land use planning, etc., and it became much more important in regions of arid and

Table 4 The trend statistics MK_Z and MMK_Z with the magnitude of the trend

ET series	MK_Z	MMK_Z	Slope Q	B
Jan	-0.70	-1.00	-0.07	85.699
Feb	-1.33	-1.33	-0.14	86.870
Mar	2.70**	2.70**	0.37	107.635
Apr	0.31	0.46	0.05	130.768
May	-0.61	-0.61	-0.14	146.940
Jun	-1.65+	-1.65+	-0.50	130.050
Jul	-2.11*	-2.11*	-0.31	101.525
Aug	-1.63	-1.63	-0.28	103.609
Sep	-2.12*	-2.12*	-0.30	121.050
Oct	-2.54*	-2.53*	-0.28	133.734
Nov	-2.73**	-2.73**	-0.20	108.000
Dec	-0.27	-0.27	-0.03	89.482
Annual	-2.81*	-2.18*	-1.49	1347.097
Winter	-1.33	-0.90	-0.20	172.420
Pre-monsoon	0.71	0.70	0.27	383.981
Monsoon	-2.60*	-2.43*	-1.03	453.393
Post-monsoon	-2.85**	-2.53*	-0.51	333.695

Significance levels: +90%; * 95%; ** 99%

semiarid zones. However, it is difficult to obtain precise field measurement due to complex nature of the process itself, spatial variability, atmospheric instability, etc. There are numerous methods of computing ET, but the Penman–Monteith (PM) equation which provides the most reliable results for the estimates of reference evapotranspiration (ET₀) in most of the regions and for varied weather condition, has been adopted by Food and Agriculture Organization (FAO). In the present study, performance evaluation and comparison of simulated ET₀ by FAO56 PM and an ANN model with single layer feed forward back propagation algorithm has been carried out. Temperature (Min., Max., and Avg.), relative humidity (Min. and Max.), average wind speed, sunshine hours and rainfall data from the meteorological monitoring station located in FRI, Dehradun on monthly basis for a period of 34 years (1980–2013) were used as inputs to predict ET₀.

The ANN models for this study were developed in neural network (nn) module of MATLAB software. The basic model architecture was optimized in term of training algorithms and number of neurons in the hidden layer of the model. Six algorithms have been tested for training the neural network, followed by evaluation of the model performance by varying the number of neurons from 1 to 10 in the hidden layer. Developed model has been evaluated using various performance indices, viz. coefficient of correlation (*r*), MSE, RMSR and NSE. Further, the trend analysis and magnitude of trend of ET₀ has been carried out using MK and MMK testing methods.

Highest value of coefficient of correlation between ANN modelled and FAO56 PM estimated ET_0 was found 0.993 and 0.956 during training and validation, respectively. Which shows the good agreement between observed and estimated ET_0 . The ANN model with “trainlm” algorithm with nine neurons was found best for simulating the ET_0 among the all iteration. In view of the ET_0 trends, the study concludes that the FRI station Dehradun is experiencing the falling trend in last 34 years at the rate of 1.49 mm/year at a significance level of 90%. The modeling approach presented in this paper demonstrated the potential of the predictive power of the ANN architecture.

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References

- Ajmera TK, Goyal MK (2012) Development of stage–discharge rating curve using model tree and neural networks: an application to Peachtree Creek in Atlanta. *Expert Syst Appl* 39(5):5702–5710
- Allen RG, Pereira LS, Raes D, Smith M (1998) Crop evapotranspiration: guidelines for computing crop requirements. FAO Irrigation and Drainage Paper No. 56, FAO Rome, Italy
- Baba APA, Shiri J, Kisi O, Fard AF, Kim S, Amini R (2013) Estimating daily reference evapotranspiration using available and estimated climatic data by adaptive neuro-fuzzy inference system (ANFIS) and artificial neural network (ANN). *Hydrol Res* 44(1):131–146
- Basistha A, Arya DS, Goel NK (2009) Analysis of historical changes in rainfall in the Indian Himalayas. *Int J Climatol* 29:555–572
- Doğan E (2008) Reference evapotranspiration estimation using adaptive neuro-fuzzy inference systems. *Irrig Drain* 58(5):617–628
- Goyal MK, Bharti B, Quilty J, Adamowski J, Pandey A (2014) Modeling of daily pan evaporation in sub tropical climates using ANN, LS-SVR, Fuzzy Logic, and ANFIS. *Expert Syst Appl* 41(11):5267–5276
- Hagan MT, Menhaj MB (1994) Training feedforward networks with the Marquardt algorithm. *IEEE Trans Neural Netw* 5(6):989–993
- Hamed KH, Rao AR (1998) A modified Mann–Kendall trend test for autocorrelated data. *J Hydrol* 204:182–196
- Hirsch RM, Slack JR, Smith RA (1982) Techniques of trend analysis for monthly water quality data. *Water Resour Res* 18(1):107–121
- Kendall MG (1962) Rank correlation methods, 3rd edn. Hafner Publishing Company, New York
- Kendall MG, Stuart A (1961) The advance theory of statistics, 2nd edn. Hafner Publishing Co, New York
- Keskin ME, Terzi O (2006) Artificial neural network models of daily pan evaporation. *J Hydrol Eng* 11(1):65–70
- Kim S, Kim HS (2008) Neural networks and genetic algorithm approach for nonlinear evaporation and evapotranspiration modeling. *J Hydrol* 351:299–317
- Kisi O (2006a) Evapotranspiration estimation using feed-forward neural networks. *Nord Hydrol* 37(3):247–260
- Kisi O (2006b) Generalized regression neural networks for evapotranspiration modelling. *Hydrol Sci J J Des Sci Hydrol* 51(6):1092–1105
- Kumar M, Raghuwanshi NS, Singh R, Wallender WW, Pruitt WO (2002) Estimating evapotranspiration using artificial neural network. *J Irrig Drain Eng ASCE* 128(4):224–233
- Kumar M, Raghuwanshi NS, Singh R (2011) Artificial neural networks approach in evapotranspiration modeling: a review. *Irrig Sci* 29(1):11–25
- Kumar D, Pandey A, Sharma N, Flügel WA (2015) Evaluation of TRMM-precipitation with rain-gauge observation using hydrological model J2000. *J Hydrol Eng* E5015007. doi:10.1061/(ASCE)HE.1943-5584.0001317
- Mann HB (1945) Nonparametric tests against trend. *Econometrica* 13:245–259
- MATLAB (Mathwork) online reference manual. <https://in.mathworks.com/help/nnet/ref/trainlm.html?requestedDomain=www.mathworks.com>. Accessed 15 Dec 2016 [online]
- Melesse AM, Ahmad S, McClain ME, Wang X, Lim YH (2011) Suspended sediment load prediction of river systems: an artificial neural network approach. *Agric Water Manag* 98(5):855–866
- Moriasi DN, Arnold JG, Van Liew MW, Bingner RL, Harmel RD, Veith TL (2007) Model evaluation guidelines for systematic quantification of accuracy in watershed simulations. *Trans ASABE* 50(3):885–900
- Odhiambo LO, Yoder RE, Yoder DC, Hines JW (2001) Optimization of fuzzy evapotranspiration model through neural training with input–output examples. *Trans ASAE* 44(6):1625–1633
- Parasuraman K, Elshorbagy A, Carey SK (2007) Modelling the dynamics of the evapotranspiration process using genetic programming. *Hydrol Sci J* 52(3):563–578
- Rahimikhoob A (2010) Estimation of evapotranspiration based on only air temperature data using artificial neural networks for a subtropical climate in Iran. *Theor Appl Climatol* 101(1–2):83–91
- Sen PK (1968) Estimates of the regression coefficient based on Kendall’s tau. *J Am Stat Assoc* 63:1379–1389
- Shamshirband S, Amirmojahedi M, Gocić M, Akib S, Petković D, Piri J, Trajkovic S (2015) Estimation of reference evapotranspiration using neural networks and cuckoo search algorithm. *J Irrig Drain Eng* 142(2):04015044
- Sudheer KP, Gosain AK, Ramasastri KS (2003) Estimating actual evapotranspiration from limited climatic data using neural computing technique. *J Irrig Drain Eng ASCE* 129(3):214–218
- Trajkovic S, Todorovic B, Stankovic M (2003) Forecasting of reference evapotranspiration by artificial neural networks. *J Irrig Drain Eng ASCE* 129(6):454–457
- Xu CY, Gong LB, Jiang T, Chen DL, Singh VP (2006) Analysis of spatial distribution and temporal trend of reference evapotranspiration and pan evaporation in Changjiang (Yangtze River) catchment. *J Hydrol* 327:81–93
- Zanetti SS, Sousa EF, Oliveira VPS, Almeida FT, Bernardo S (2007) Estimating evapotranspiration using artificial neural network and minimum climatological data. *J Irrig Drain Eng ASCE* 133(2):83–89