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Abstract: Contamination from agricultural sources and, in particular, nitrate pollution, is one of the main concerns in groundwater management. However, this type of pollution entails the entrance of other substances into the aquifer, as well as it may promote other processes. In this study, we deal with hydrochemical and isotopic analysis of groundwater samples from four distinct zones in Catalonia (NE Spain), which include different lithological units, to investigate the influence of manure fertilization on the overall hydrochemical composition of groundwater. Results indicate that high nitrate concentrations, resulting from intense manure application, homogenize the contents of the major dissolved ions (i.e.; Cl-, SO42-, Ca2+, Na+, K+, and Mg2+). Moreover, positive linear relationships between nitrate and some ions are found indicating the magnitude of the fertilization impact on groundwater hydrochemistry. Nevertheless, the increasing concentration of specific ions is not only attributed to the manure input, but to the enhancing effect of manure and slurry upon the biogeochemical processes that control water-rock interactions. Such results raise awareness that such processes should be evaluated in advance in order to assess adequate groundwater resources assessment.

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Dear Sir,

We submit the manuscript entitled "Nitrate pollution of groundwater; all right..., but nothing else?", by A. Menció, J. Mas-Pla, N. Otero, O. Regàs, M. Boy-Roura, A. Folch, R. Puig, C. Domenech and J.Bach, to be considered for publication in *Science of the Total Environment*.

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Yours sincerely,

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Girona, May 11, 2015.



Highlights

- The effects of nitrate pollution have been evaluated in five different aquifer types
- Statistical and multivariate analyses are used to identify groundwater changes
- Agricultural pollution modifies groundwater conditions and geochemical processes

1 2 3 4 5 6	1	Nitrate pollution of groundwater; all right, but nothing else?
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• The effects of nitrate pollution have been evaluated in five different aquifer types

- Statistical and multivariate analyses are used to identify groundwater changes
- Agricultural pollution modifies groundwater conditions and geochemical processes

INTRODUCTION

Nitrate occurrence and transport in aquifers have been widely studied since it is one of the major threats in groundwater, and most aquifers in agricultural areas are affected by this contaminant (Spalding and Exner, 1993; EEA, 2012). Most of the papers that focused on nitrate pollution study, by means of hydrochemical and isotopic data, nitrate trends in groundwater at different scales to identify potential sources of pollution and build-up hydrogeochemical models to understand the behavior of nitrate polluted aquifers (for instance, Burg and Heaton, 1998; Hudak, 2000; Katz et al., 2004; Masetti et al., 2008; Vitòria et al., 2008; Carbó et al., 2009; Kaown, et al., 2009; Menció et al., 2011; Boy-Roura et al., 2013). However, agricultural pollution due to a long and continued application of organic fertilizer (slurry and manure) may not only entail an increase on this ion in aquifers. On the one hand, other substances present in manure also enter into the groundwater system, such as

50 ions, metals, emerging organic contaminants, or even microorganisms. On the other hand, 51 these substances interact with the subsurface environment, modifying groundwater 52 conditions, enhancing geochemical processes and even modifying groundwater communities 53 (Cho et al., 2000; Böhlke, 2002; Murray et al., 2010; Stein et al., 2010; Korbel and Hose, 54 2011; Lapworth et al., 2012; Korbel et al., 2013; Choi et al., 2013).

In this study, we look for the evidence of the changes that manure application exerts on the major components concentrations by conducting an analysis of a large hydrochemical database, including isotopic data, of nitrate polluted aquifers representative of several geological environments. Our aim is to analyze whether manure and slurry application lead to a loss of the hydrogeological fingerprint of the geological background, and whether the resulting groundwater composition is only caused by mixing with manure lixiviates or by changes in biogeochemical processes.

Manure applied as fertilizer has a complex chemical composition. For instance, Vitòria (2004) analyzed its composition from several farms in Osona (Catalonia, NE Spain) and reported that the concentration of some major inorganic components in pig manure, the most used organic fertilizer in the studied areas in Catalonia (NE Spain), is notably larger than their natural values in groundwater. This is the reason why its impact on groundwater hydrochemistry must be evaluated to avoid misinterpretation of supposedly hydrochemical natural data.

Samples for this study belong to four distinct regions in Catalonia, all of them classified as Nitrate Vulnerable Zones (NVZs) as a result of the transposition of the Nitrate Directive (ND) 91/676/EC. In fact, NVZs in Catalonia cover up to 40% of the total area, and half of them show nitrate concentrations higher than 40 mg/L, affecting 17 out of the 53 groundwater bodies which are at risk of not meeting the European Water Framework Directive goals (ACA, 2007; Boy-Roura, 2013). In this study, we analyze datasets from the

following NVZs, according to published results in the Selva basin (Folch et al., 2011; Menció et al., 2012; Puig et al., 2013), Empordà basin (Puig, 2014), Osona region (Vitòria et al., 2008; Otero et al., 2009; Menció et al., 2011; Boy-Roura et al., 2013), and Garrotxa area (based on still unpublished data). According to their lithologies, five groups of aquifers are distinguished (Figure 1 and Figure 2):

- Group 1: Aquifers in igneous rocks, especially granite and granodiorite and, in a lesser degree, in metamorphic rocks such as shale, schist, marble, and gneiss. These aquifers are mainly located in the ranges surrounding the Selva and Empordà basins. Hydrochemically, they present Ca²⁺-HCO₃⁻ and Ca²⁺-HCO₃⁻-Cl⁻ facies, tending to evolve to Na⁺-HCO₃⁻.
- Group 2: Aquifers in sedimentary rocks, mainly Paleogene sedimentary rocks, including detritic, organic and chemical sedimentary rocks, such as conglomerates, sandstones, siltstones, limestones, marls, and gypsum deposits. Although these aquifers are present in all the study areas, they are mainly exploited in Osona and Garrotxa, and with a minor extent in Empordà. Main hydrochemical facies are Ca²⁺-HCO₃⁻, Ca²⁺-Mg²⁺-HCO₃⁻ or Ca²⁺-SO₄²⁻.
- Group 3: Aquifers in sediments derived from igneous and metamorphic rocks. The arkosic sands, gravels and silt layers of the Neogene deposit that filled the Empordà and Selva areas, and the more recent Quaternary alluvial formations associated to the main rivers in these basins, are the weathering products of the main range areas, mainly located in the Pyrenees Montseny-Guilleries and Gavarres ranges. They show similar facies than Group 1.

• Group 4: Aquifers in sediments derived from sedimentary rocks. These aquifers are located in Osona and Garrotxa basins, and constitute the Quaternary alluvial aquifers and surface formations in these areas. Their facies are analogous to those of Group 2.

Group 5: Aquifers in volcanic materials. In this group of aquifers there is a broad variety of materials, from phreatomagmatic and pyroclastic deposits to basaltic lava flow locally interleaved by sedimentary (alluvial) levels, which may locally be the most productive units. These materials are mainly located in the Garrotxa area, but volcanic lithologies also constitute aquifers in some parts of the Selva basin. This group shows the lowest EC values with Ca^{2+} -HCO₃⁻ or Ca^{2+} -Mg²⁺-HCO₃⁻ facies.

METHODOLOGY

Field surveys, in the studied datasets (Folch et al., 2011; Menció et al., 2012; Puig et al., 2013; Puig, 2014; Otero et al., 2009; Menció et al., 2011; Boy-Roura et al., 2013), were conducted from 2006 to 2013. A total number of 204 groundwater sampling locations constitute the whole dataset -distributed as Selva basin: 37 samples; Empordà basin: 45; Osona region: 57; Garrotxa area: 65- and they are classified according to the five described lithological groups (Table 1).

Similar analytical procedures for hydrochemical and isotopic analysis were followed at each studied region. Specific details are described in each of the given references. Concentration units are reported in mg/L. Most of the samples had an ionic mass balance error between \pm 5%. Isotope δ -notation for water isotopes is expressed in terms of the % deviation of the isotope ratio of the sample relative to that of the V-SMOW standard. Analytical errors are $\pm 0.06\%$ for δ^{18} O and $\pm 0.7\%$ for δ D. Nitrate isotope notation is also expressed in terms of δ (‰) relative to that of the international standards AIR (atmospheric N₂) for δ^{15} N. Precision $(\equiv 1\sigma)$ of the samples are $\pm 0.3\%$ and $\pm 0.4\%$, for $\delta^{15}N_{NO3}$ and $\delta^{18}O_{NO3}$, respectively.

The SPSS program (version 19, 2010, SPSS Inc.) was used to conduct the different statistical analyses. Differences between aquifer hydrochemical parameters were analyzed considering the Kruskal-Wallis and Mann-Whitney U tests, for non-parametric data, since none of the

parameters were normally distributed. In addition, a Principal Component Analysis (PCA) was conducted in order to identify associations between variables and samples. The associations obtained through a PCA, based on similar magnitudes and variations in chemical, physical, and isotopic values in the groundwater samples composition, are adequate to indicate the influence of human factors, hydrochemical processes, or even, the origin of groundwater (for instance, Helena et al., 2000; Menció and Mas-Pla, 2008; Menció et al., 2013; Re et al., 2014).

RESULTS AND DISCUSSION

At a first glance, hydrochemical characteristics of groundwater samples with low NO_3^{-1} content (below the health standard limit of 50 mg NO₃^{-/}L) reflect the dominant role of lithology in the composition, as observed in Figure 2. Mean and standard deviation values (Table 1) and boxplot shapes of selected parameters (EC, Cl⁻, SO₄²⁻, Na⁺, Ca²⁺; Figure 2) for samples with low nitrate content, reflect the differences among aquifer types attributed to distinct lithology, showing the hydrochemical characteristics of groundwater samples under natural conditions. For instance, aquifers in groups 1 and 3, including igneous and metamorphic aquifers and sediments derived from these rocks, do not show significant differences between concentrations of the major hydrochemical components (with p-values between 0.089 and 0.881). These aquifers, for instance, present significant differences when compared with groups 2 and 4, representing aquifers in sedimentary rocks and in sediments derived from these rocks, for ions such as Cl⁻, Na⁺, and Ca²⁺ (p-values ranging from values <0.001 to 0.038). Contrarily, hydrochemical characteristics of groundwater at high-nitrate concentrations (> 50 mg NO_3/L ; Figure 2) seem to homogenize their concentrations despite lithological differences, presenting wider ranges for major ions distribution and erasing the

main differences in their hydrochemical composition of the distinct groups observed at low
NO₃⁻ samples.

In detail, when samples of low and high nitrate content for each lithological aquifer group are compared, significant higher EC values are observed in the most polluted wells (p-values ranging from 0.006 to values lower than 0.001). Higher EC values are related to significant higher concentrations of ions: Cl^- , SO_4^{2-} , Na^+ and Ca^{2+} (Table 1). An exception are samples of group 1, which do not show significant differences in any of these parameters between low and high nitrate groups (p-values ranging from 0.123 for Ca^{2+} to 0.877 for Cl⁻), since some of these samples belong to thermal and/or CO2-rich systems. Geochemical evolution in such environments overlays the chemical contribution of manure inputs. In addition, it is worth mentioning that these Cl⁻ and Na⁺ higher concentrations are not detected in wells located in aquifer group 3 (in sediments derived from igneous and metamorphic rocks, and p-values of 0.684 and 0.844, respectively), where natural processes can also favor high Cl^{-} and Na^{+} concentrations.

It is also relevant that fewer hydrochemical differences are detected when high nitrate concentration sample populations for the different aquifers are compared (Figure 2 and Table 1). As expected due to lithological similarity, no significant differences are detected between groups 1 and 3 (p-values ranging from 0.087 for SO₄²⁻ and 0.906 for Cl⁻), neither between groups 2 and 4 (p-values between 0.050 for HCO₃⁻ and 0.732 for SO₄²⁻). Nevertheless, when samples with high nitrate concentrations of group 1 and 2 are compared, only significant differences in SO42- and K⁺ are detected (p-values of 0.024 and 0.023), while other components show similar concentration ranges. Furthermore, differences among high nitrate concentrations samples of groups 1 and 4 are only evident for Ca²⁺ and K⁺ (p-values of 0.018 and 0.011); comparing samples of groups 2 and 3, significant differences are reduced to Mg^{2+} and K^+ (with p-values <0.001); and, differences between samples of groups 1 and 5 with high

nitrate concentrations are significant for Na⁺ and K⁺ (p-values of 0.014 in both cases). Indeed, this group 5 shows the largest differences with the rest of the aquifer groups for high NO_3^{-1} samples. In volcanic materials, nitrate pollution is clearly lower than in the rest of the studied aquifers, with an overall mean value of 30.90±3.81 mg/L (Table 1). In this case, high nitrate concentration samples presented differences with respect to nitrate for samples of groups 2, 3 and 4 (p-values of 0.026, 0.027 and 0.031, respectively). Between group 5 and these other groups, significant differences were also observed in EC, Cl⁻, Ca²⁺ and SO₄²⁻ (p-values ranging from <0.001 to 0.044); besides, significant differences were detected between group 5 and groups 2 and 3 for Ca^{2+} (p-values <0.001); and, with group 3 for K⁺ and Mg²⁺ (p-values of 0.008 and 0.004, respectively).

Complementarily, two Principal Component Analyses were conducted to determine the different relationships among samples and variables. The first one considered all available samples, a total number of 204; and considered eleven variables, including: EC, pH, total aqueous concentration of HCO₃⁻, Cl⁻, SO₄²⁻, Ca²⁺, Mg²⁺, Na⁺ and K⁺, and δ^{18} O and δ D. In the second analysis, the dataset included only those samples that, in addition to the mentioned variables, also included nitrate isotopic data ($\delta^{15}N_{NO3}$). In this analysis the number of samples was reduced to 158, as in 32% of them (mainly from group 5) lacked the nitrate isotopic information. By conducting both statistical analyses we seek illustrating whether denitrification processes are relevant on the understanding of the hydrochemical changes in groundwater related to nitrate pollution. This is the reason why both analyses are shown herein, and compared.

In the first PCA, variables as pH, HCO_3^- , K^+ and δD were ruled out of in order to obtain better values of the goodness-of-fit statistics. Thus, a PCA conducted with the 7 remaining variables had a Barlett chi-square statistic of 1241.6 (for 28 degrees of freedom and a minimum significance level of <0.001), and a value of sampling adequacy (MSA) obtained

by the Kaiser-Meyer-Olkin of 0.732. In the second analysis, the final solution considered 8
variables, and had a better fit with a Barlett chi-square statistic of 1085.5 (for 36 degrees of
freedom and a minimum significance lower than 0.001), and a MSA of 0.758.

Four varifactors (VF) were obtained in both analyses, explaining 91.59% and 88.19% of the total variance, respectively. In each PCA, these varifactors explained the same processes and associations of variables; excepting natural attenuation, which was only considered in the second one as $\delta^{15}N_{NO3}$. For this reason, only the scores and samples distribution obtained in the second PCA are shown in Table 2 and Figure 3. Thus, after a Varimax rotation, the final VFs were interpreted as follows:

 VF1 includes SO₄²⁻, Mg²⁺, Ca²⁺, and EC, explaining a 32.53% of the total variance. This VF1 stands for water-rock interaction processes, specifically in sedimentary formations where gypsum dissolution occurs, since VF1 highest values are obtained in samples located in aquifers belonging to groups 2 and 4, where these processes have been described to govern groundwater hydrochemistry (Figure 3a; Menció et al., 2011; Soler et al., 2014).

• VF2 represents the 20.41% of variability of the data. It mainly includes NO₃⁻, with a lower participation of Ca²⁺, EC and Cl⁻. VF2 has been interpreted as nitrate pollution, showing the direct relationship between NO₃⁻ and EC, Ca²⁺ and Cl⁻ variables.

With a 20.23% of the total variance, VF3 is participated by δ¹⁸O, Na⁺ and Cl⁻. This VF3 is linked to the recharge altitude of groundwater samples, with high scores in samples recharged at low altitude, and low scores in samples recharged at high altitude. In addition, it is worth recalling that recharge areas of most of the wells drilled in igneous and metamorphic rocks and in the sediments derived from them (groups 1 and 3) are located at low altitude. Because of their lithology, water samples

present high concentrations of Na⁺ and Cl⁻, explaining the association of both ions with δ^{18} O (Table 1, Figure 2 and Figure 3).

• VF4 is mainly associated to $\delta^{15}N_{NO3}$, with a lower participation of Na⁺ and Cl⁻. With a 15.03% of the total variance explained, this last VF represents natural attenuation processes (i.e., denitrification). The association of Na⁺ and Cl⁻ with high $\delta^{15}N_{NO3}$ values ($\delta^{15}N>15\%$, according to Kendall et al., 2007) may be attributed to the attenuation processes described in the Selva basin, group 3, where regional flow systems with Na⁺-HCO₃⁻ facies and reducing conditions, enhanced natural heterotrophic attenuation processes (Puig et al., 2013).

Two sample tendencies with positive scores for SO_4^{2-} , Ca^{2+} and Mg^{2+} (that is, VF1) are distinguished in a plot VF1 vs. VF2. The first tendency presents high scores for VF2 as well, which links samples with high SO_4^{2-} , Ca^{2+} and Mg^{2+} content to high nitrate concentrations (quadrant I in Figure 3a), including samples from aquifer groups 2, 3 and 4 with nitrate concentrations between 280 to 590 mg NO_3 /L. The second one is associated to negative values in VF2, and it is composed of samples from aquifer groups 2 and 4, those related to sedimentary rocks and the sediments derived from these materials. In this case, nitrate concentrations are notably lower, between 2.5 to 119 mg/L, and the sample with the highest VF1 score is the one with the lowest nitrate concentration. Thus, two clear sets of samples with distinct origins for SO_4^{2-} are distinguished in these study areas: one with SO_4^{2-} related to nitrate pollution, and the other with SO₄²⁻ related to natural sources.

In Figure 3b, high and low scores for nitrate pollution (VF2) are distributed in both high and low altitude recharge areas (as defined by δ^{18} O isotopic compositions, VF3). This indicates a widespread nitrate distribution all over the study areas, being consistent with an intense agricultural activity. However, most of samples in aquifer group 5, which wells are located in volcanic rocks, present low VF2 values, consistent with the comparatively low pollutionlevels in this aquifer type (Table 1).

Finally, when nitrate content (VF2) and nitrate attenuation (VF4) are plotted (Figure 3c), samples with the highest nitrate concentration (with values exceeding 150 mg NO₃⁻/L at VF2 scores > 1) are located in quadrants I and II; while quadrants I and IV include those samples affected by natural attenuation (with δ^{15} N>15‰). Samples of quadrant I and IV indicate that attenuation processes are active independently of the nitrate concentration. Besides, in some cases, nitrate content has been reduced to values lower than the health standard limits of 50 mg/L, as represented by the sample with negative scores of VF2. These nitrate attenuation processes have been detected in all aquifers types.

Bivariate plots of selected ions against nitrate concentrations point out their linear increase proportional to manure application. SO_4^{2-} and Ca^{2+} (both included in VF1) and Cl⁻ (participating in VF2 and VF3) as major components illustrate this fact (Figure 4). Linear regression equations are estimated for each aquifer lithology using all the available data, despite their $\delta^{15}N_{NO3}$ value. Such regression equations prevail over two facts: 1) the inherent variability of the manure chemical composition depending on its origin and storage before application; and 2) the heterogeneity of soil processes, which are different, in type and magnitude, depending on the soil nature and the crop type. Therefore, equations indicate the rate of ion concentration changes with increasing nitrate content (slope); that is, the amount of each ion that any unit of nitrate adds to groundwater whether as a direct input or as a result of enhancing geochemical processes; and the expected value of the major component where no fertilization occurs (y-intercept). Such linear increase is quite evident for these three ions, and similar relationships occur for other major components.

Nevertheless, each geological environment show distinct responses to nitrate inputs. Waterrock interaction in igneous and metamorphic rocks (group 1) will depend on the low

solubility of silicate minerals, and of other processes as cation exchange, which may significantly alter the $Ca^{2+}:Na^{+}$ ratio and it may be responsible for a wide dispersion of data values with and without the occurrence of nitrate pollution. The large heterogeneity of volcanic deposits (group 5), which may include sedimentary layers among volcanic episodes, also originates a wide range of geochemical values as regards of SO_4^{2-} and Ca^{2+} . However, Cl⁻ concentration in the studied aquifers appears to be well related to manure fertilization. Moreover, the occurrence of thermal and/or CO₂-rich waters in group 1 aquifers may enhance specific weathering processes resulting, for instance, in large chloride concentrations that infiltrating water may dilute.

As already suggested by the PCA results, SO_4^{2-} content may originate from evaporitic rocks (gypsum) dissolution or from manure contribution. These two geochemical paths are also recognized in groups 2 and 4, where sedimentary rocks or their denudation deposits constitute the aquifers: significantly large values of SO_4^{2-} are found at low nitrate concentration (<75 mg NO_3/L , whereas a clear linear relationship between them appears as nitrate content increases. Correlation factors, as they appear in Figure 4, are misleading since they represent the whole dataset; exclusion of those points largely affected by gypsum dissolution will turn to larger R^2 values proving the influence of manure fertilization on the final groundwater composition.

290 Denitrification processes also affect such bivariate relationships. In those places where 291 autotrophic nitrate reduction has been proved (group 2 samples, Osona region; Otero et al., 2009), an increase of SO_4^{2-} due to pyrite oxidation will occur coupled to nitrate decreease; 293 whereas heterotrophic attenuation processes will increase alkalinity. Both processes will also 294 have consequences on the gypsum and calcite equilibriums modifying the Ca²⁺ concentration, 295 yet the neat linear increase of Ca²⁺ with nitrate in most of the aquifer types might suggest that 296 it directly originates in the manure itself. However, a geochemical insight to calcium

equilibrium indicates that nitrification of reduced N from manure generates acidity along with NO_3^- and it enhances carbonate dissolution. This process can even be accentuated where lime or dolomite are applied as soil neutralizers in agricultural lands (Böhlke, 2002; Choi et al., 2013), increasing Ca^{2+} and Mg^{2+} concentrations. Complementarily, the analysis of hydrochemical evolution of pig manure stored in experimental pits showed a decrease of Ca²⁺ and Mg²⁺ through time (Vitòria, 2004). Since manure is usually stored for several months before its application, Ca^{2+} increase shown in Figure 4 may be better attributed to geochemical reactions occurring in the soil and in the subsurface following fertilization than to direct manure contribution. Hence, the observed hydrochemical composition is not just a simple conservative mixture of groundwater with infiltrating slurry and manure lixiviation. Contrarily, Na⁺ and K⁺ showed a conservative behavior in the liquid manure, being their concentrations increased by evaporation (Vitòria, 2004). However, groundwater samples are not specifically rich in both cations (Figure 2), suggesting that other processes as plant uptake, sorption and cation exchange may control their final content in groundwater (Böhlke, 2002).

CONCLUSIONS

Data from aquifers with distinct lithological environments prove that manure application as fertilizer modifies groundwater hydrochemical composition, turning it notably distinct than the expected natural background given by water-rock interaction and other natural processes. Manure and slurry fertilization homogenize the overall hydrochemistry despite lithological differences, hindering the geochemical interpretation inherent to any regional groundwater resources evaluation study.

Nevertheless, such compositional modifications which, in general, tend to adopt a linear increase (Figure 4, all data) cannot be solely attributed to the effect of direct manure inputs, 322 as fertilization may enhance or reduce geochemical processes that control groundwater 323 composition at equilibrium. Therefore accepting that nitrate pollution influences, 324 groundwater regional reports should look for the relationships shown in these aquifers; 325 checking potential influences upon the expected (natural) hydrochemical composition, so a 326 better understanding of the hydrogeological system is attained.

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443 FIGURE CAPTIONS

Figure 1: Geographical and geological setting of the distinct study areas.

Figure 2. Box plots of the main hydrochemical parameters according to the aquifer groups and nitrate concentrations ranges: Low nitrate concentrations refer to values below 50 mg/L, and high nitrate concentrations refer to values above 50 mg/L (see SI2 for details). Legend: Boxes represent the 25 and 75 percentiles and the median; while a white line represents the mean value. Bars define the 10 and 90% percentile, and dots refer to extreme values.

Figure 3. PCA scores distribution of the different samples, considering denitrification, 451 plotted according the lithological group and the $\delta^{15}N_{NO3}$, content: a) VF1 vs VF2, b) VF3 vs 452 VF2, and c) VF4 vs VF2.

Figure 4. Bivariate relationships of $SO_4^{2^-}$, Ca^{2^+} and Cl^- vs. NO_3^- . Linear regression equations are calculated for all the samples of the group, disregarding their $\delta^{15}N$ value. *Legend*: red circles, samples with $\delta^{15}N > 15\%$; black circles, samples with $\delta^{15}N < 15\%$; white circles, samples with $\delta^{15}N$ data not available.

TABLES

Table 1: Mean values and standard errors of the main physicochemical characteristics of the

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different aquifer groups, differentiating samples of low and high nitrate content.

	G1- Igneous and metamorphic rocks		G2- Sedimentary rocks		G3- Aquifers derived from igneous and metamorphic rocks		G4- Aquifers derived from sedimentary rocks		G5- Volcanic materials	
	<50 mg/L	>50 mg/L	<50 mg/L	>50 mg/L	<50 mg/L	>50 mg/L	<50 mg/L	>50 mg/L	<50 mg/L	>50 mg/L
Number of wells	9	4	28	36	15	27	22	29	29	5
EC (µS/cm)	819±110	915±135	819±70	1124±65	841±72	1073±78	770±27	1070±64	623±21	779±31
рН	7.18±0.14	6.99±0.04	7.35±0.09	7.43±0.06	7.47±0.12	7.28±0.08	7.18±0.04	7.08±0.04	7.47±0.06	7.25±0.21
Eh (mV)	236.7±62.1	386.5±39.3	190.1±20.1	247.4±27.4	354.9±38.7	379.8±8.3	161.0±13.5	164.3±17.3	274.6±21.0	240.0±10.0
T (°C)	17.4±0.5	16.9±0.3	15.4±0.6	15.9±0.4	16.3±0.4	16.4±0.2	15.1±0.4	13.2±0.2	14.5±0.6	14.5±0.4
O ₂ (mg/L)	1.5±0.6	3.9±1.4	4.0±0.6	4.0±0.5	3.7±1.1	5.2±0.5	4.6±0.6	4.4±0.4	8.0±0.6	7.6±1.3
HCO3 ⁻ (mg/L)	368.1±59.6	346.9±40.0	418.3±15.4	388.0±13.0	355.1±34.4	360.7±13.5	411.9±18.6	407.4±13.1	284.8±13.0	341.7±39.0
Cl ⁻ (mg/L)	101.0±24.7	99.2±42.8	42.4±6.7	94.9±10.4	86.1±15.1	91.8±11.9	16.0±1.6	73.3±12.4	16.1±2.2	30.7±10.4
SO_4^{2-} (mg/L)	35.9±5.7	57.2±8.6	132.4±46.4	164.3±32.1	68.4±16.8	98.1±12.4	43.5±5.4	184.6±34.6	48.0±6.5	41.4±8.2
NO3 ⁻ (mg/L)	13.7±4.2	105.3±20.3	17.0±2.5	166.3±19.4	18.6±4.4	137.5±16.6	25.4±2.7	172.2±25.2	23.9±2.4	71.4±10.0
Na ⁺ (mg/L)	104.2±38.0	79.7±18.6	34.8±4.2	54.3±7.5	63.5±12.0	56.4±5.1	12.5±1.4	42.0±5.6	16.5±1.5	21.5±.9
K ⁺ (mg/L)	2.2±0.4	1.0±0.3	4.4±1.7	8.5±2.4	2.4±0.4	5.8±3.0	2.8±0.6	8.8±2.1	7.6±0.7	6.9±1.7
Ca ²⁺ (mg/L)	82.6±14.0	127.5±21.5	132.6±16.2	185.7±10.1	110.7±12.7	166.3±9.7	135.8±7.1	196.0±10.8	87.2±4.5	89.5±8.4
Mg ²⁺ (mg/L)	14.6±3.6	22.0±3.5	37.4±5.4	38.2±3.8	19.3±2.0	22.7±2.4	21.0±1.1	44.1±5.6	19.5±1.2	34.1±3.5

2 3	466	parameters.		
4 5				VF1 VF2 VF3 VF4
6 7			$\overline{\mathrm{SO_4}^{2-}}$ (mg/L)	0.965 0.008 -0.005 0.039
8 9			Mg^{2+} (mg/L)	0.885 0.181 0.007 0.249
11			Ca^{2+} (mg/L)	0.740 0.562 0.004 0.035
13			EC (µS/cm)	<u>0.641</u> 0.546 0.425 0.184
14			NO_3^- (mg/L)	0.121 0.949 0.112 0.057
17			$\delta^{18}O_{\rm H2O}$ (‰)	-0.238 0.097 0.848 -0.178
19			Na^+ (mg/L)	0.295 0.054 0.738 0.426
20			Cl ⁻ (mg/L)	0.289 0.520 0.601 0.405
22			δ^{15} N _{NO3} (‰)	0.113 0.085 0.047 0.935
24 25			Eigenvalue	2.928 1.837 1.821 1.352
26 27			% Variance	32 53 20 41 20 23 15 03
28 29	467			
30 31				
32 33	468			
34 35	469			
36 37	470			
38 39	470			
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Table 2. Loadings obtained in the second PCA, including $\delta^{15}N_{NO3}$, among the different 466 parameters.





Aquifer groups



Figure 4 Click here to download high resolution image



