



Measurement report: Rapid decline of aerosol absorption coefficient and aerosol optical property effects on radiative forcing in an urban area of Beijing from 2018 to 2021

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Abstract. Reliable observations of aerosol optical properties are crucial for quantifying the radiative forcing of climate. The simultaneous measurements of aerosol optical properties at three wavelengths for PM₁ and PM₁₀ were conducted in urban Beijing from March 2018 to February 2022. The aerosol absorption coefficient (σ_{ab}) at 550 nm of PM₁₀ and PM₁ decreased by 55.0 % and 53.5 % from 2018 to 2021. The significant reduction in σ_{ab} may be related to reduced primary emissions caused by effective air pollution control measures. PM_{2.5} mass concentration decreased by 34.4 % from 2018 to 2021. Single scattering albedo (SSA) increased from 0.89 ± 0.04 for PM₁₀ (0.87 ± 0.05 for PM₁) in 2018 to 0.93 ± 0.03 for PM₁₀ (0.91 ± 0.04 for PM₁) in 2021. Increasing SSA and decreasing PM_{2.5} mass concentration suggest that the fraction of absorbing aerosols decreased with improved air quality due to pollution control measures being taken. The annual average submicron absorption ratio (R_{ab}) increased from 86.1 % in 2018 to 89.2 % in 2021, suggesting that fine particles are the main contributors to total PM₁₀ absorption and that the contribution of fine particles to absorption became more important. The absorption Ångström exponent (AAE) in winter decreased from 2018 to 2021, implying a decreasing contribution from brown carbon to light absorption, which may relate to the reduced emissions of biomass burning and coal combustion. During the study period, aerosol radiative forcing efficiency became more negative, mainly influenced by increasing SSA and was -27.0 and -26.2 W m^{-2} per aerosol optical depth (AOD) for PM₁₀ and PM₁ in 2021. Higher σ_{ab} and PM_{2.5} mass concentrations were primarily distributed in clusters 4 and 5, transported from the south and the west of Beijing each year. σ_{ab} and PM_{2.5} corresponding to clusters 4 and 5 decreased evidently from 2018 to 2021, which may result from the control of source emissions in surrounding regions of Beijing. The 4-year data presented in this study provide critical optical parameters for radiative forcing assessment within two size ranges and are helpful for evaluating the effectiveness of clean air action.

1 Introduction

Atmospheric aerosols perturb the Earth's atmospheric radiation balance and climate forcing by directly affecting the scattering and absorption of solar radiation (Charlson et al., 1992; Jacobson, 2001) but also by indirectly affecting cloud reflectivity and precipitation processes (Twomey, 2007). Light scattering aerosols contribute to offsetting the warming effect of CO₂, while absorbing aerosols contribute to the heating of the atmosphere (Bond and Bergstrom, 2007) and produce a positive radiative forcing (Segura et al., 2016). The largest contribution to aerosol absorption is from black carbon (BC), which absorbs strongly over the entire solar spectrum (Bond and Bergstrom, 2007). Dust and brown carbon (BrC) are also light absorption aerosols, which strongly absorb in the ultraviolet (UV) spectrum. Globally, aerosols contributed an effective radiative forcing (ERF) of $-1.3 \pm 0.7 \text{ W m}^{-2}$, and the ERF due to emissions of BC is now estimated to have been $0.11 (-0.20 \text{ to } 0.42) \text{ W m}^{-2}$ between 1750 to 2019 (Szopa et al., 2021). However, aerosol properties are highly spatially and temporally variable, which results in radiative forcing variation from local to global scales and creates an observational challenge (Collaud Coen et al., 2013; Ealo et al., 2018; Andrews et al., 2011). Therefore, reliable observations of aerosol optical properties are crucial for quantifying the radiative forcing of climate.

In order to assess the role of aerosols on climate forcing accurately, a set of parameters that describe aerosol's optical properties are needed, such as scattering coefficient (σ_{sp}), absorption coefficient (σ_{ab}), backscatter fraction (b), and single scattering albedo (SSA). The SSA is a key variable that determines the magnitude and the sign of the aerosol forcing (Hansen et al., 1997; Lee et al., 2007; J. Li et al., 2022; Zhang et al., 2020). Previous studies found that SSA values range from slightly less than 0.8 to almost purely scattering particles with SSA close to 1 at locations worldwide (Laj et al., 2020; Pandolfi et al., 2018), and higher SSA values indicate a tendency towards a cooling effect (J. Li et al., 2022). b describes how much aerosol particles scatter radiation in the backward hemisphere compared with the total scattering, which is a crucial variable for aerosol radiative forcing efficiency (RFE) calculations (Andrews et al., 2011; Sheridan and Ogren, 1999; Luoma et al., 2019). Previous studies found that the magnitude of the RFE increases with increasing b (Shen et al., 2018). Typical values of b for the atmospheric aerosol at 550 nm were from approximately 0.05 to 0.20 (Titos et al., 2021).

Moreover, aerosol optical properties are wavelength-dependent, and the absorption Ångström exponent (AAE) describes the spectral dependence of light absorption by aerosols and is typically used to differentiate between different aerosol types (Helin et al., 2021). The AAE for fresh BC is ~ 1 , indicating “weak” spectral dependence of light absorption (Bond et al., 2013; Bond and Bergstrom, 2007), and the AAE > 1 indicates the presence of BrC or dust, which

tends to exhibit absorption that increases sharply as the wavelength decreases (Moosmüller et al., 2009; Lack and Cappa, 2010). Thus, obtaining the aerosol absorption coefficient at different wavelengths is essential and can be helpful to differentiate between different aerosol types.

As one of the world's most populous and rapidly developing megacities, Beijing experienced rapid economic growth and urbanization, accompanied by severe air pollution. Many in situ measurements of aerosol optical properties have been conducted in Beijing (Bergin et al., 2001; He et al., 2009; Garland et al., 2009; Jing et al., 2015; Wang et al., 2019; Zhao et al., 2019; Xia et al., 2020). Previous studies found that high aerosol loading leads to large σ_{ab} in Beijing (Jing et al., 2015; Garland et al., 2009; Bergin et al., 2001). Moreover, the AAE showed significant seasonal variations in Beijing. The significantly higher AAE in winter than in summer highlights the important role of the absorption of non-BC components (e.g. BrC) in winter (Xie et al., 2020; Xia et al., 2020). In order to reduce emissions and improve air quality, the government implemented strict pollution control measures (Xu and Zhang, 2020). Significant decreases in PM_{2.5} mass concentrations were found in Beijing, and the annual mean elemental carbon (EC) concentrations declined from 4.0 to 2.6 $\mu\text{g m}^{-3}$ from March 2013 to February 2018 in Beijing (Ji et al., 2019). Xia et al. (2020) separated and quantified the effects of emission control and meteorological transport variability on BC loading from 2015 to 2017 in the North China Plain. However, the environmental effects caused by emission controls are related to not only their mass concentrations but also their optical properties and radiative effect (Luo et al., 2020). Therefore, it is necessary to investigate the multi-year variations in aerosol optical properties and radiative effect in providing a comprehensive understanding of the effects of emission control. Wang et al. (2019) found that the absorption coefficient (σ_{ap}) for PM_{2.5} decreased from 2014 to 2017, with a significant decrease in σ_{ap} in autumn. Sun et al. (2022) estimated that the direct radiative forcing of BC decreased by 67 % from $+3.36 \text{ W m}^{-2}$ in 2012 to $+1.09 \text{ W m}^{-2}$ in 2020. However, these studies were mostly conducted for total suspended particulate matter, PM_{2.5} size cut, or PM₁₀ size cut. Few studies focused on the submicron and supermicron particle optical properties and estimated aerosol radiative effect in the era after the “Action Plan on Prevention and Control of Air Pollution”. Acquiring the aerosol optical for the total ($< 10 \mu\text{m}$ diameter) and submicron aerosol is also in line with the recommendation of the aerosol advisory group of the Global Atmosphere Watch (WMO/GAW, 2016).

In this study, the simultaneous measurements of aerosol optical properties at three wavelengths for PM₁ and PM₁₀ were conducted in urban Beijing from March 2018 to February 2022. The annual, seasonal, and diurnal variations in aerosol optical properties for two size cuts were investigated. The scattering properties of aerosols for two size ranges (PM₁₀ and PM₁) under dry conditions observed in

Beijing have been analysed in detail by Hu et al. (2021). Thus, this study mainly focused on the variation in aerosol absorption coefficient, single scattering albedo, and absorption Ångström exponent for PM₁₀ and PM₁. Moreover, the aerosol radiative effects in two size cuts were estimated. Finally, the transport and its impact on aerosol optical properties were analysed. The 4-year data presented in this study provide key optical parameters for radiative forcing assessment within two size ranges and are helpful for evaluating the effectiveness of clean air action.

2 Instrumentation and methods

2.1 Site description

The sampling site in this study is located on the roof of the Chinese Academy of Meteorological Sciences (CAMS; 39°57' N, 116°19' E; 46 m a.s.l. – above sea level) in Beijing, which is a typical urban site in the northwest of Beijing between the 2nd and 3rd ring roads. The laboratory is on the roof of the CAMS building, and the measurements are taken at 53 m a.g.l. (above ground level). The site is mainly influenced by local emissions from residential living and traffic pollution (Xia et al., 2019).

2.2 Instruments and measurements

The ambient air was sampled into a PM₁₀ impactor at 16.7 L min⁻¹ and then to an adsorption aerosol dryer, which controlled the relative humidity (RH) of sample air to below 30 % (Tuch et al., 2009). The dried aerosol sample passes through switched impactors that toggle the aerosol size cut between 1.0 µm (< 1 µm) and 10 µm (< 10 µm) aerodynamic particle diameters every 30 min, thus allowing us to measure both fine and coarse particles (Hu et al., 2021). The sample aerosol was then passed into the nephelometer (TSI Inc., Model 3563) and tricolor absorption photometer (TAP; Brechtel Manufacturing, Inc., Hayward, CA, USA).

The integrating nephelometer measured the scattering coefficient (σ_{sp}) (angular range of 7–170°) and backscattering coefficient (σ_{bsp}) (angular range of 90–170°) at 450, 550, and 700 nm. The scattering and backscattering coefficient were corrected for truncation and instrument non-idealities using the method described by Anderson and Ogren (1998). Details are given in Hu et al. (2021). To ensure the data's accuracy and reliability, the nephelometer was calibrated regularly using filtered ambient air using a HEPA filter and CO₂ with a purity of 99.999 %. A zero-check was automatically performed once per hour to obtain a nephelometer background.

TAP measures the absorption coefficient (σ_{ab}) at 465, 520, and 640 nm with the 47 mm diameter, glass-fibre filter and is a commercially available version of the continuous light absorption photometer (CLAP), which is low cost and has high sensitivity (Ogren et al., 2017). The TAP comprises eight sample spots and two reference spots. The aerosol-laden air

passes through one sample spot at a time, which allows for 8 times the filter lifetime compared to single-spot photometers (Davies et al., 2019). Unlike the multi-angle absorption photometer (MAAP), TAP does not require a co-located aerosol light scattering or extinction measurement to derive aerosol light absorption (Ogren et al., 2017). Thus, simultaneous observation of aerosol light scattering has been measured and used to correct absorption data. When the nephelometer and TAP were calibrated or malfunctioning, no data are available. During this study, 84 % of the data were effective.

2.3 Data processing

The TAP measures the light transmitted through a filter as particles are deposited onto the filter. The filter attenuation coefficient (σ_{atn}), at a specific wavelength (λ), can be determined as

$$\sigma_{\text{atn}}(\lambda) = \frac{A}{Q} \times \frac{\Delta \text{atn}(\lambda)}{\Delta t}, \quad (1)$$

where $\Delta \text{atn}(\lambda)$ is the filter attenuation at times t_1 and t_2 , A is the area of the filter, and Q is the sample flow rate through the filter.

In order to correct the error caused by multiple scattering and filter loading, the aerosol light absorption coefficient ($\sigma_{\text{ab}}(\lambda)$) was corrected based on the methods of Bond et al. (1999) and Ogren et al. (2017). First, the effect of filter loading was calibrated based on Eq. (2):

$$\sigma_{\text{ab}}(\lambda)_{\text{raw}} = \frac{0.85 \times \sigma_{\text{atn}}(\lambda)}{K_2 \times (1.0796 \times \text{Tr}(\lambda) + 0.71)}. \quad (2)$$

Then, $\sigma_{\text{ab}}(\lambda)_{\text{raw}}$ at 465, 520, and 640 nm was adjusted to the wavelength of the light scattering coefficient based on the calculated AAE. Finally, the multiple scattering effect was corrected based on Eq. (3):

$$\sigma_{\text{ab}}(\lambda) = \sigma_{\text{ab}}(\lambda)_{\text{raw}} - \frac{K_1 \times \sigma_{\text{sp}}(\lambda)}{K_2}, \quad (3)$$

where $\text{Tr}(\lambda)$ is the normalized filter transmittance at time t relative to transmittance at the start of sampling ($t = 0$) and σ_{sp} is the aerosol light scattering coefficient at 450, 550, and 700 nm measured by the nephelometer. K_1 and K_2 were derived by Bond et al. (1999) as $K_1 = 0.02 \pm 0.02$ and $K_2 = 1.22 \pm 0.20$, where the uncertainties are given for the 95 % confidence level.

Using the corrected absorption coefficient data, the following parameters were calculated.

The absorption Ångström exponent (AAE) describes the spectral dependence of light absorption.

$$\text{AAE} = - \frac{\ln \left(\frac{\sigma_{\text{ab}}^{\lambda_1}}{\sigma_{\text{ab}}^{\lambda_2}} \right) / \left(\frac{\sigma_{\text{ab}}^{\lambda_2}}{\sigma_{\text{ab}}^{\lambda_1}} \right)}{\ln(\lambda_1/\lambda_2)} \quad (4)$$

The submicron absorption ratio (Rab) is determined as the ratio of the absorption coefficients for PM₁ and PM₁₀.

$$\text{Rab} = \frac{\sigma_{\text{ab}}(D < 1 \mu\text{m})}{\sigma_{\text{ab}}(D < 10 \mu\text{m})}, \quad (5)$$

where σ_{ab} ($D < 1 \mu\text{m}$) and σ_{ab} ($D < 10 \mu\text{m}$) are σ_{ab} for particle diameters < 1 and $10 \mu\text{m}$, respectively.

Aerosol radiative forcing efficiency (RFE) at the top of the atmosphere (TOA) is a simplified formula that describes how large of an impact the aerosols would have on the aerosol radiative forcing (ΔF) per unit of aerosol optical depth (AOD) (Sheridan and Ogren, 1999), and we estimated the RFE at TOA as Eq. (6) (Haywood and Shine, 1995; Sheridan and Ogren, 1999):

$$\text{RFE} = \frac{\Delta F}{\text{AOD}} = -DS_0T_{\text{at}}^2(1 - A_c) \times \text{SSA} \times \beta \times (1 - R_s)^2 - \left(\frac{2R_s}{\beta}\right) \times \left(\frac{1}{\text{SSA}} - 1\right), \quad (6)$$

where D is the fractional day length, S_0 is the solar constant, T_{at} is the atmospheric transmission, A_c is the fractional cloud amount, and R_s is the surface reflectance. The constants used were $D = 0.5$, $S_0 = 1370 \text{ W m}^{-2}$, $T_{\text{at}} = 0.76$, $A_c = 0.6$, and $R_s = 0.15$ as suggested by Haywood and Shine (1995), and upper scatter fraction β was calculated from $\beta = 0.0817 + 1.8495 \times b - 2.9682 \times b^2$. The backscatter fraction (b) was calculated based on scattering coefficient (σ_{sp}) and backscattering coefficient (σ_{bsp}) measured by the nephelometer as $b = \sigma_{\text{bsp}}/\sigma_{\text{sp}}$. Equation (4) has been widely used to assess the intrinsic radiative forcing efficiency of aerosols at the TOA (Sheridan and Ogren, 1999; Virkkula et al., 2011; Shen et al., 2018). Note that the RFE in this study was in a dry condition. As the backscatter fraction and single scattering albedo are all RH-dependent, the RFE is also sensitive to RH (Fierz-Schmidhauser et al., 2010). Previous studies revealed that the RFE increased as the elevating RH (Titos et al., 2021; Xia et al., 2023). In this study, the values of ΔF at TOA were also calculated by multiplying the RFE for PM_{10} with the AOD of ambient atmospheric aerosols observed at the CAMS site during the study period. AOD can be downloaded from the Aerosol Robotic Network (AERONET). Note that the RFE was in a dry state; thus the ΔF at TOA here may be slightly underestimated.

2.4 Other data used

The hourly $\text{PM}_{2.5}$ and PM_{10} mass concentrations were measured at Guanyuan station, which is about 3 km from the CAMS site. The data can be derived from the national air quality real-time publishing platform (<http://www.cnemc.cn>, last access: 11 May 2023). The hourly meteorological data were measured at Haidian station (station no. 54399) and obtained from the National Meteorological Information Center of the China Meteorological Administration.

2.5 Back trajectories analysis

To investigate the influence of air mass origins on aerosol optical properties, 48 h backward trajectories arriving at Beijing at a height of 500 m a.g.l. were calculated from 00:00 to

23:00 LST (local standard time) each day from March 2018 to February 2022, using the TrajStat software, combined with the HYSPLIT 4 model (Hybrid Single-Particle Lagrangian Integrated Trajectory), and the National Centers for Environmental Prediction (NCEP) Global Data Assimilation System (GDAS) data with a $1^\circ \times 1^\circ$ resolution (Draxler and Hess, 1998; Wang et al., 2009).

In this study, the four seasons are defined as follows: spring from March to May, summer from June to August, autumn from September to November, and winter from December to the following February, and all data are reported in Beijing time (UTC+8).

3 Results and discussion

3.1 Temporal variation in aerosol optical properties

Figure 1 shows the annual variation in σ_{ab} , SSA, R_{ab} , and $\text{PM}_{2.5}$ mass concentration from 2018 to 2021. During the study period, the annual mean $\text{PM}_{2.5}$ in 2018 was $54.7 \mu\text{g m}^{-3}$, and it decreased by 34.4 % ($35.9 \mu\text{g m}^{-3}$) in 2021, which suggested that the strict pollution control measures are effective in reducing the PM loadings in Beijing (Lei et al., 2021). Gong et al. (2022) demonstrated that emission reduction dominated the variations in $\text{PM}_{2.5}$ mass concentration in Beijing from 2013 to 2020, and meteorology and emission reduction contributed 7 % and 63.2 % of decreases, respectively. σ_{ab} at 550 nm of PM_{10} and PM_1 showed similar annual variations. The annual mean σ_{ab} at 550 nm of PM_{10} and PM_1 decreased by 55.0 % and 53.5 %, respectively, from 2018 to 2021. The Mann–Kendall trend test supported the results that the decrease in σ_{ab} for PM_1 and PM_{10} from 2018 to 2021 was significant (Table S1 in the Supplement). Carbonaceous aerosol, especially black carbon, is closely related to aerosol absorption (Yang et al., 2009). A continuous decrease in σ_{ab} was consistent with the continuous reduction in black carbon concentration observed in Beijing in previous studies (Ji et al., 2019; Sun et al., 2022), which was mainly related to significantly reduced primary emissions caused by effective air pollution control measures in recent years (Xia et al., 2020). The annual mean σ_{ab} for PM_{10} and PM_1 in 2021 was 9.8 and 8.7 Mm^{-1} , which were both lower than the result observed in Nainital, in the Gangetic Himalayan region, India (Dumka et al., 2015), and the measurement at an urban site in Spain from March 2006 to February 2007 (Titos et al., 2012). In fact, with the emission reduction and improvement in air quality, the aerosol scattering coefficient (σ_{sp}) for PM_{10} and PM_1 also decreased in Beijing. Hu et al. (2021) revealed that σ_{sp} decreased by approximately 18.4 % for PM_{10} and 16.7 % for PM_1 from 2018 to 2019 in Beijing. Atmospheric conditions also have an effect on aerosol optical properties. The variations in meteorological parameters from 2018 to 2021 (Fig. S4) showed that pressure, wind speed, temperature, and RH varied slightly, while accumulated precipitation increased in 2021 compared

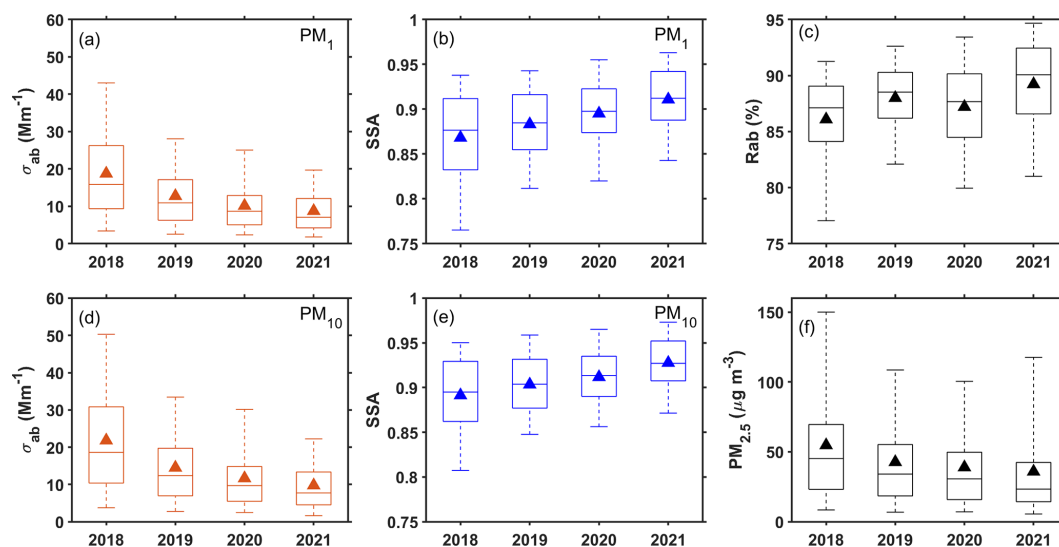


Figure 1. Annual variation in aerosol optical properties and $\text{PM}_{2.5}$ mass concentration: absorption coefficient σ_{ab} at 550 nm for (a) PM_1 and (d) PM_{10} , SSA at 550 nm for (b) PM_1 and (e) PM_{10} , (c) Rab, and (f) $\text{PM}_{2.5}$ mass concentration. The solid line inside the box represents the median, and the triangle indicates the mean. The box contains the range of values from 25 % (bottom) to 75 % (top), and the upper and lower whiskers are the 95th and 5th percentiles, respectively.

with the other 3 years. On the other hand, a correlation analysis was made between aerosol optical properties and meteorological parameters. The Pearson correlation coefficients (R) between σ_{ab} and meteorological parameters (Table S2) are lower than 0.5, indicating that a weak correlation ($R < 0.5$) was found between σ_{ab} and meteorological parameters. This suggests that the meteorological parameters' influence on σ_{ab} is minor. Xia et al. (2020) revealed that the effect of emission reduction was the major reason for the decrease in BC in Beijing. Actually, the reduction trend of σ_{ab} was observed at a background station in China, as well as at European stations, which have time series longer than 10 years. σ_{ab} showed a statistically significant decreasing trend at Mt. Waliguan, a background station in China, from 2008–2018 (Collaud Coen et al., 2020), which was similar to a decreasing trend of black carbon (BC) at Mt. Waliguan from 2008–2017, mainly related to emission reduction (Dai et al., 2021). A statistically significant decrease in 10-year σ_{ap} was found in 12 stations in Europe, which was similar to a decreasing trend in BC concentration in Europe related primarily to traffic emission decreases (Collaud Coen et al., 2020).

SSA is a key variable in assessing the aerosol radiative forcing. The variation in SSA also reflects the ratio of aerosol scattering to total extinction with aerosol composition changes. The annual variations in SSA for PM_{10} and PM_1 were similar. During 2018–2021, annual mean SSA at 550 nm increased from 0.89 ± 0.04 for PM_{10} (0.87 ± 0.05 for PM_1) in 2018 to 0.93 ± 0.03 for PM_{10} (0.91 ± 0.04 for PM_1) in 2021. Increasing SSA and decreasing $\text{PM}_{2.5}$ mass concentration during the past 4 years suggested that the fraction

of absorbing aerosols became lower compared to scattering aerosols with the improvement in air quality due to pollution control measures being taken. Collaud Coen et al. (2020) found that SSA observed at Mt. Waliguan, a background station in Asia, presented an increasing trend based on 10-year datasets, which were related to more recent abatement policies. The mean submicron absorption ratio (Rab) increased yearly during the same period. It was from 86.1 % in 2018 to 89.2 % in 2021, suggesting that fine particles are the main contributors to total PM_{10} absorption, and the contributions from fine particles to absorption became more important.

The σ_{ab} , SSA, and AAE for PM_1 and PM_{10} showed similar annual variations in all seasons (Figs. 2 and S1 in the Supplement). Thus, if not stated otherwise, the following discussion takes the aerosol optical properties of PM_{10} as an example. As shown in Fig. 2 seasonal averages of σ_{ab} presented a continuous reduction during all seasons from 2018 to 2021, reflecting the reduction in absorbing aerosols which were related to effective control of absorbing aerosol emissions in Beijing. σ_{ab} decreased by half in autumn and winter during the study period, which was probably due to reducing coal consumption as a heating source and the reduction in biomass burning. Compared with 2018, σ_{ab} in the winters of 2019, 2020, and 2021 decreased by 3.0 %, 24.9 %, and 53.2 %, respectively. In the winter of 2019, the lockdown of COVID-19 caused emission reduction from human activities in China (Le et al., 2020; Tian et al., 2020); however, the unexpected smallest reduction in σ_{ab} was observed in the winter of 2019 compared with the winters of 2020 and 2021. This is related to the fact that severe haze pollution still occurred in the North China Plain, and BC concentrations rose unex-

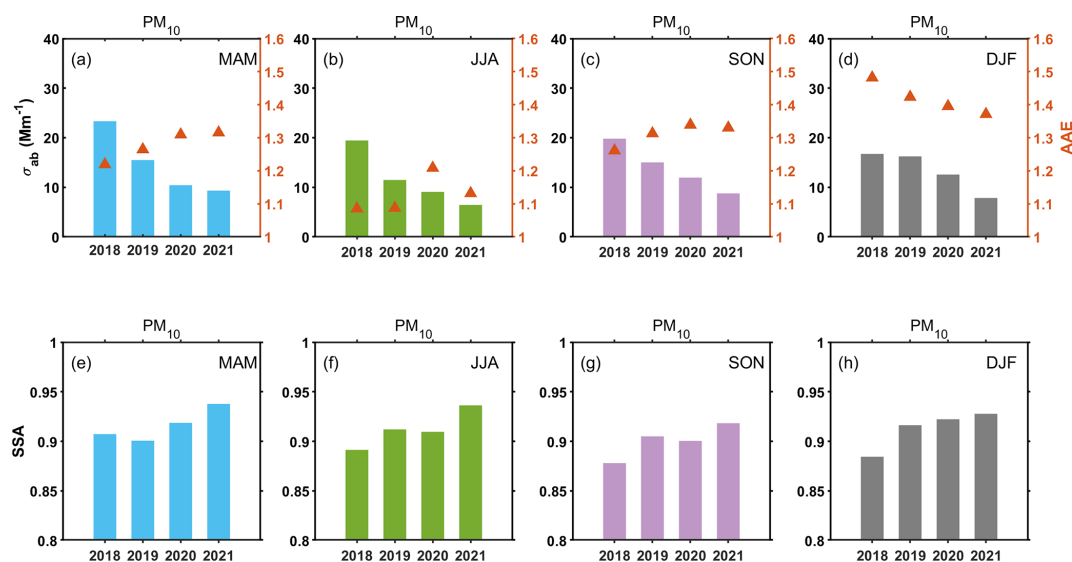


Figure 2. Seasonal variation in aerosol optical properties of PM_{10} from 2018–2021: (a–d) σ_{ab} (bar) at 550 nm and $\text{AAE}_{450/700}$ (triangle) and (e–h) SSA (bar) at 550 nm.

pectedly during the lockdown period (Liu et al., 2021; Jia et al., 2021). In particular, σ_{ab} for PM_1 and PM_{10} decreased even up to 63 % and 67 % in the summer from 2018 to 2021. Traffic is a relatively stable source of absorption aerosols in summer (W. Li et al., 2022). The largest deduction of σ_{ab} was in summer and could be related to more strict vehicle emission standards (Zhang et al., 2019).

In general, the AAE was lowest in summer and highest in winter. The mean values of the AAE for PM_{10} were 1.13 and 1.41 in summer and winter, respectively, similar to results at an urban site in Beijing in 2018 (Xie et al., 2020). During summer, the average AAE was generally close to 1, which suggested that BC from traffic emissions was the major component of light-absorbing aerosols. W. Li et al. (2022) found that the percentage of liquid fuel (traffic) contributing to the total BC was 86.8 % in summer in Beijing. The highest AAE suggested that BrC contributed to light absorption strongest in winter, which is due to enhanced emissions from biomass burning and coal combustion in winter (Sun et al., 2018). Notably, the AAE decreased in winter from 1.48 for PM_{10} (1.48 for PM_1) in 2018 to 1.37 for PM_{10} (1.34 for PM_1) in 2021 (Figs. 2 and S1), indicating a decreasing contribution from BrC to light absorption, which may relate to the effect the control of biomass burning and coal combustion had by changes in the heating energy structure (Ji et al., 2022). To improve air quality, the Beijing–Tianjin–Hebei region adjusted the energy structure during the heating period and developed clean heating projects, such as the “coal to gas” project (Zhao et al., 2020; Liu et al., 2019). During the whole period, the AAE was similar in spring and autumn, indicating that light-absorbing aerosols were from similar emission sources in spring and autumn (Ran et al., 2016). The AAE slightly increased in spring and autumn from 2018

to 2021. Part of the reason was the occurrence of multiple fugitive dust events in spring and autumn (Yi et al., 2021; Gui et al., 2022). On the other hand, BrC could also be formed from secondary reactions (Bond et al., 2013; Wang et al., 2022). A slight increase in the AAE in spring and autumn may also have been caused by a greater amount of secondary organic aerosol formation as a result of an increased atmospheric oxidation capacity (Ji et al., 2019; Lei et al., 2021).

The seasonal mean SSA increased in all seasons from 2018 to 2021, indicating that the contribution of scattering aerosols to extinction increased. This suggested that more effective control of scattering aerosols should have more importance attached to it in order to improve visibility in the future. In particular, SSA in winter increased significantly from 0.88 in 2018 to 0.93 in 2021, which revealed that the proportion of absorbing aerosols decreases considerably in winter. This is consistent with recent research which suggests that air pollution control measures have been more effective in reducing the primary pollution emissions than secondary species (Vu et al., 2019; Sun et al., 2020). On the other hand, seasonal mean SSA for PM_{10} was 0.94 ± 0.04 , 0.94 ± 0.04 , 0.92 ± 0.04 , and 0.93 ± 0.03 in spring, summer, autumn, and winter 2021. Similar SSA suggests that the proportions of light absorbing and scattering components became relatively stable in the four seasons.

Figure 3 shows the diurnal variations in σ_{ab} and SSA at 550 nm for PM_{10} , which are similar to those for PM_1 (Fig. S2). In the past 4 years, σ_{ab} was lower during the day and higher at night in the four seasons. This was consistent with that observed at an urban site in Beijing during 2014–2016 (Wang et al., 2019). The evolution of the planetary boundary layer had an important influence on the diurnal variation in the σ_{ab} . With stronger solar radiation, the

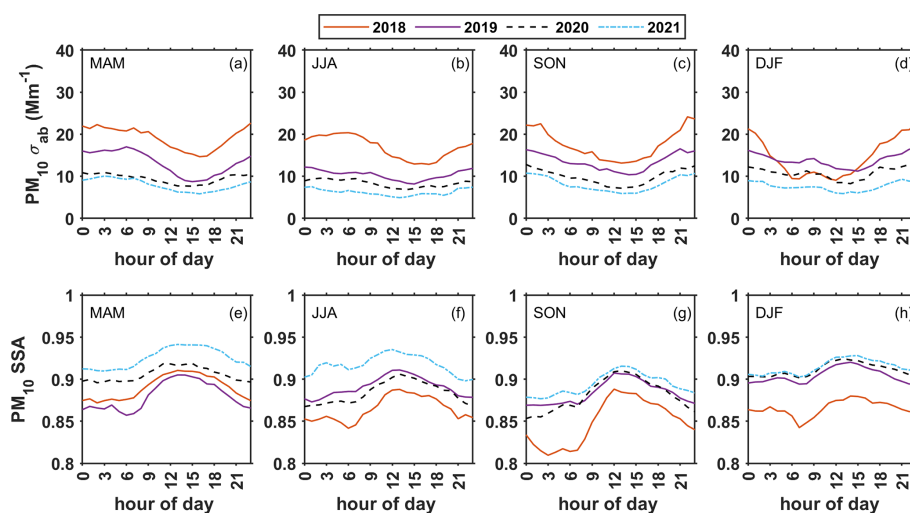


Figure 3. Diurnal variations in σ_{ab} (a–d) and SSA (e–h) at 550 nm for PM_{10} in the four seasons from 2018 to 2021.

boundary layer was more fully developed during the daytime, and after sunset, the convective boundary layer underwent a transition to the nocturnal stable boundary layer (Guo et al., 2016). Furthermore, emissions also affected the diurnal variation in the σ_{ab} . For example, heavy-duty diesel trucks and heavy-duty vehicles were only allowed to enter urban areas from 23:00 LST to the following day at 06:00 LST (Hu et al., 2021). As a response, the minimum σ_{ab} occurred during 12:00–18:00 LST, when the planetary boundary layer was well-developed and truck emissions were lower. With shallow boundary layer height and enhanced emissions from heavy-duty trucks, σ_{ab} reached the maximum at night. During the study period, SSA showed a significant peak in the afternoon in the four seasons, which was similar to previous studies in urban Beijing (Zhao et al., 2019; Wang et al., 2019). Higher SSA was shown in the afternoon, which was mainly related to the reduction in absorbing aerosol emissions and more secondary scattering aerosol produced by strong chemical reactions under intensive solar radiation and high temperature in the afternoon (Han et al., 2017).

3.2 Aerosol radiative effect

To study the climate impact of the aerosol particles, we investigated the variation in aerosol RFE at the TOA. As seen in Fig. 4, RFEs for PM_{10} and PM_1 were always negative during the whole observation period, suggesting that the aerosols measured in urban Beijing have a stable cooling effect on the climate. RFEs for PM_{10} and PM_1 under dry conditions were -27.0 and -26.2 W m^{-2} per AOD in 2021 in urban Beijing, which was slightly more negative than that of -24.9 W m^{-2} per AOD in Nanjing (Shen et al., 2018) and significantly more negative than that of -19.9 W m^{-2} per aerosol optical depth in Finland (Virkkula et al., 2011). This suggested that the aerosols in urban Beijing have a higher cooling effi-

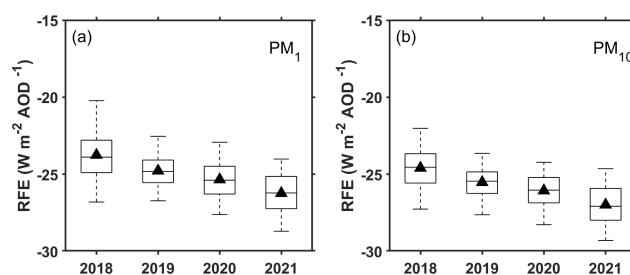


Figure 4. The annual variation in aerosol radiative forcing efficiency for PM_1 (a) and PM_{10} (b). The solid line inside the box represents the median, and the triangle indicates the mean. The box contains the range of values from 25 % (bottom) to 75 % (top), and the upper and lower whiskers are the 95th and 5th percentiles, respectively.

ciency. In Eq. (6) the fractional day length (D), solar constant (S_0), atmospheric transmission (T_{at}), fractional cloud amount (A_c), and surface reflectance (R_s) were constants, which were widely used in previous studies (Delene and Ogren, 2002; Andrews et al., 2011; Sherman et al., 2015; Shen et al., 2018). These values are the globally averaged values and do not always represent the conditions in Beijing, but using the same constants makes it possible to compare the intrinsic forcing efficiencies of the aerosols measured at different stations around the world and to study how the RFE changes with varying SSA and b (Sherman et al., 2015; Luoma et al., 2019). On the other hand, the RFE is sensitive to RH, as the aerosol optical properties are different due to hygroscopic growth (Fierz-Schmidhauser et al., 2010; Luoma et al., 2019). Previous studies demonstrate that SSA increases with RH, while b decreases with increasing RH (Carrico et al., 2003; Cheng et al., 2008). The change in SSA to increase with RH and in b to decrease with RH will have opposite effects on the RFE, and thus to some extent the RH dependen-

cies of these two parameters will counterbalance each other (Luoma et al., 2019). Titos et al. (2021) found that the range of forcing enhancement in different types of sites varies from almost no enhancement up to a factor of 3–4 at RH = 90 %. The results observed in urban Beijing showed that the aerosol radiative forcing at RH = 80 % was 1.48 times that under dry conditions (Xia et al., 2023). The RFE was calculated in a dry state in this study, while the atmosphere is not generally dry in the ambient air. Thus, the RFE in this study does not represent ambient conditions. The simplified RFE in this study does not represent the actual value for the aerosol forcing; however, it can still indicate how the changes in aerosol optical properties affect the climate (Delene and Ogren, 2002; Andrews et al., 2011; Sherman et al., 2015). The RFE was affected by SSA and b , and we investigated the RFE variations with SSA and b in Beijing. As shown in Fig. 5, when SSA increases from 0.7 to 0.92, the mean RFE increases by 1.59 times, suggesting that SSA plays an important role in strengthening cooling efficiency. When SSA > 0.92, the mean RFE keeps relatively constant. The approximate constant RFE does not mean that the absolute aerosol radiative forcing is constant; it just suggests that the intrinsic nature of the aerosol will not significantly affect the calculation of the RFE (Andrews et al., 2011). Also, the backscatter fraction has a negative relationship with the RFE. A lower value of backscatter fraction corresponds to larger particles (Luoma et al., 2019). The RFE became more negative with increasing b , suggesting that smaller particles would cool the atmosphere more efficiently. During the study period, SSA increased from 0.89 to 0.93, while the yearly mean value of b was 0.13 every year during the study period. The RFE became more negative from 2018 to 2021, suggesting that the efficiency of the aerosol cooling in the atmosphere was higher, which was mainly influenced by increasing SSA.

The ratio of ΔF / AOD is known as the aerosol RFE, and ΔF at the TOA was calculated by multiplying the RFE for PM₁₀ with the AOD of ambient atmospheric aerosols observed at the CAMS site during the study period. As shown in Fig. 6, the mean value of ΔF from 2018–2021 was -15.0 , -12.5 , -12.1 , and -11.8 W m⁻², respectively. Although the RFE became more negative, the annual mean ΔF in 2021 corresponding to lower columnar aerosol loading became less negative than that of 2018, corresponding to higher columnar aerosol loading (Fig. S3), which was consistent with the analysis that aerosol loading was an essential factor for the estimation of ΔF (Andrews et al., 2011; Delene and Ogren, 2002).

3.3 Transport and its impact on aerosol optical properties in Beijing

In addition to local emissions, regional transport is also an important source of particulate matter in Beijing (Chang et al., 2019). Based on previous studies, aerosol source regions and air mass pathways could also affect aerosol optical prop-

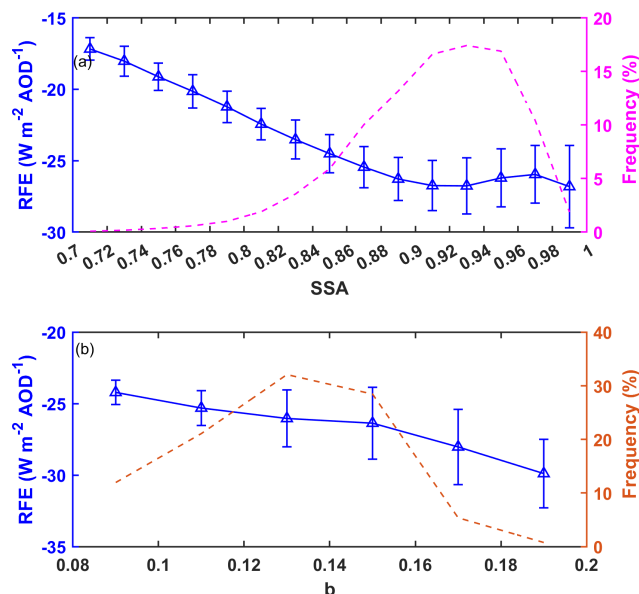


Figure 5. The relationship of RFE with (a) SSA and (b) backscatter fraction. The dashed pink line represents the frequency distribution of SSA (a), and the dashed brown line represents the frequency distribution of the backscatter fraction (b).

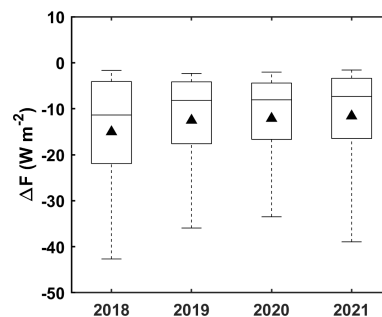


Figure 6. Annual variation in aerosol radiative forcing (ΔF) at the TOA from 2018 to 2021 calculated from daily mean data. The solid line inside the box represents the median, and the triangle indicates the mean. The box contains the range of values from 25 % (bottom) to 75 % (top), and the upper and lower whiskers are the 95th and 5th percentiles, respectively.

erties, and the different origins of air masses showed different aerosol optical properties (Zhuang et al., 2015; Pu et al., 2015). The air mass back-trajectory analysis in the North China Plain revealed that the absorption coefficients and SSA were high when the air masses came from densely populated and highly industrial areas (Yan et al., 2008). Therefore, air mass back trajectories were analysed in this study to explore the regional transports' influence on aerosol optical properties. First, the air mass back trajectories during 2018–2021 were calculated and clustered (Fig. 7); then, we studied the statistics on the aerosol optical properties of each cluster from 2018–2021 (Fig. 8). Based on the Euclidean

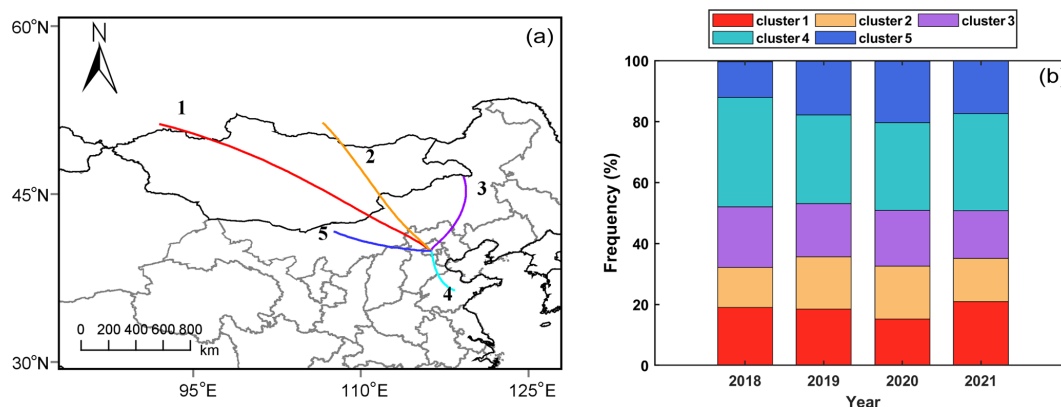


Figure 7. (a) Air mass clusters of back trajectories arriving in Beijing during 2018–2021 and (b) the fraction of each cluster accounting for the total back trajectories in each year.

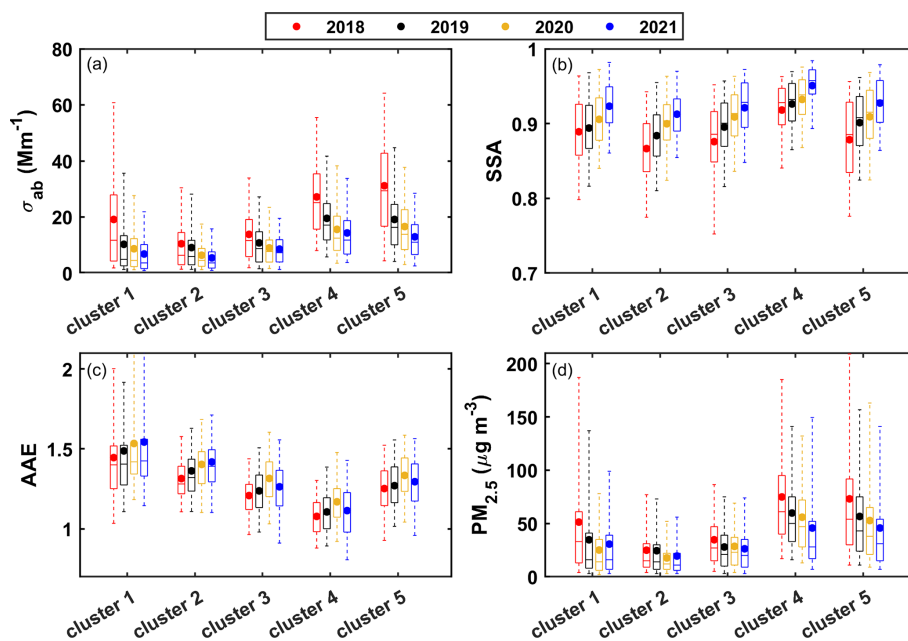


Figure 8. The variation in (a) σ_{ab} , (b) SSA, (c) AAE, and (d) $PM_{2.5}$ mass concentration in each cluster from 2018 to 2021. The solid line inside the box represents the median, and the dot indicates the mean. The box contains the range of values from 25 % (bottom) to 75 % (top), and the upper and lower whiskers are the 95th and 5th percentiles, respectively.

distance, the back trajectories were classified into five clusters, in which clusters 1, 2, and 3, which originated from the clean areas in Mongolia and eastern Inner Mongolia and were transported to Beijing along the pathway with low emissions, corresponded to low σ_{ab} and low $PM_{2.5}$ (Fig. 8a and d). Cluster 4 from the south of Beijing and cluster 5 from the west of Beijing were referred to as the polluted air masses, and the average $PM_{2.5}$ concentrations and σ_{ab} of clusters 4 and 5 were higher than those of clusters 1, 2, and 3 in each year (Fig. 8a and d). Cluster 4 passed through Shandong and Hebei provinces and was heavily polluted before arriving in Beijing. Cluster 5 passed through polluted Shanxi and Hebei during transport. Higher σ_{ab} and $PM_{2.5}$ mass concentrations

were mainly distributed in clusters 4 and 5 each year. Lower AAE in cluster 4 indicates that the southern air mass carries more freshly emitted BC particles. The SSA of cluster 4 from the south was higher (Fig. 8b), which may relate to low BC / $PM_{2.5}$ ratios in southern air masses (Xia et al., 2020). Zhang et al. (2013) found that high levels of secondary inorganic aerosols related to high humidity were transported by southern air masses, which enhanced heterogeneous reactions and led to relatively low BC / $PM_{2.5}$ ratios. Figure 7b shows the percentage of each cluster, accounting for the total back trajectories in each year. The results indicated that the variation in each cluster fraction from 2018 to 2021 was slight. In general, clusters 1–5 accounted for 19%–21%,

13 %–17 %, 16 %–20 %, 29 %–36 %, and 12 %–20 % of total back trajectories, respectively. Notably, the percentage of pollution-relevant air masses (clusters 4 and 5) was ~ 50 % each year, indicating that the transport from the south and the west has a considerable impact on the aerosol optical properties. σ_{ab} corresponding to clusters 4 and 5 decreased by 47.3 % and 58.4 %, and a decrease in $\text{PM}_{2.5}$ mass concentration from clusters 4 and 5 was 38.9 % and 37.4 % during 2018–2021 (Fig. 8a and d), which may result from the air quality improvement caused by the control of source emissions in surrounding regions of Beijing. Therefore, the comprehensive control of atmospheric pollution in Beijing and surrounding regions would be highly effective in reducing air pollution in Beijing.

4 Conclusions

In this study, 4-year measurements of aerosol absorption properties and single scattering albedo for PM_{10} and PM_1 in Beijing were analysed. The annual mean $\text{PM}_{2.5}$ in 2018 was $54.7 \mu\text{g m}^{-3}$, and it decreased by 34.4 % ($35.9 \mu\text{g m}^{-3}$) in 2021, which suggested that the strict pollution control measures are effective in reducing the PM loadings in Beijing. The annual mean σ_{ab} of PM_{10} and PM_1 decreased by 55.0 % and 53.5 %, respectively, and it showed a similar decrease in all seasons. The significant reduction in σ_{ab} may be related to reduced primary emissions caused by effective air pollution control measures. The SSA at 550 nm increased from 0.89 ± 0.04 for PM_{10} (0.87 ± 0.05 for PM_1) in 2018 to 0.93 ± 0.03 for PM_{10} (0.91 ± 0.04 for PM_1) in 2021, and the seasonal averages of SSA for the two sizes also increased in the four seasons. Increasing SSA and decreasing $\text{PM}_{2.5}$ mass concentration suggest that the fraction of absorbing aerosols decreased with improved air quality due to pollution control measures being taken. During the study period, the annual average of R_{ab} increased year by year and was up to 89.2 % in 2021, indicating that fine particles are the main contributors to the total PM_{10} particle absorption and that the contributions from fine particles to absorption became more important in Beijing.

During the study period, the AAE was lowest in summer and highest in winter. The seasonal mean AAE in summer was generally close to 1, indicating that freshly emitted BC from traffic sources was a major component of light-absorbing aerosols. The highest AAE highlights the importance of BrC light absorption in winter. Notably, the AAE in winter decreased from 2018 to 2021, implying a decreasing contribution from BrC to absorption, which may relate to the effective control of biomass burning and coal combustion caused by changes in the heating energy structure. The AAE in spring and autumn was similar, indicating light-absorbing aerosols were from similar emission sources in these two seasons.

Using a simple analytical equation, we investigated the aerosol radiative effect. Aerosol RFE for PM_{10} and PM_1 was always negative, suggesting that the aerosols measured in urban Beijing have a stable cooling effect on the climate. RFEs for PM_{10} and PM_1 under dry conditions were -27.0 and -26.2 W m^{-2} per aerosol optical depth in 2021 in urban Beijing. The RFE was influenced by SSA and b . Higher b corresponds to more negative RFE, suggesting that smaller particles would cool the atmosphere more efficiently. When $\text{SSA} < 0.92$, the absolute value of mean RFE increased by 1.59 times, suggesting that SSA plays an important role in strengthening cooling efficiency. When $\text{SSA} > 0.92$, the mean RFE keeps relatively constant, suggesting that the intrinsic nature of the aerosol will not significantly affect the calculation of the RFE. The SSA increased from 0.89 to 0.93, while the yearly mean value of b was 0.13 every year during the study period. The RFE became more negative from 2018 to 2021, suggesting that the efficiency of the aerosol cooling in the atmosphere was higher, which was mainly influenced by increasing SSA.

Regional transport and its impact on aerosol optical properties were also analysed. The air mass back trajectories arriving at Beijing were divided into five clusters. Clusters 1, 2, and 3, which originated from the clean area in Mongolia and eastern Inner Mongolia, were transported to Beijing along the pathway with low emissions, corresponding to low σ_{ab} and low $\text{PM}_{2.5}$. Air masses from the south and west (clusters 4 and 5), which both crossed the polluted region, always brought high $\text{PM}_{2.5}$ concentrations and σ_{ab} . σ_{ab} corresponding to clusters 4 and 5 decreased by 47.3 % and 58.4 %, and a decrease in $\text{PM}_{2.5}$ mass concentration from clusters 4 and 5 was 38.9 % and 37.4 % during 2018–2021, which may result from the control of source emissions in surrounding regions of Beijing. Therefore, a comprehensive control of atmospheric pollution in surrounding regions of Beijing is conducive to reducing pollution in Beijing.

Data availability. The data in this study are available at <https://doi.org/10.5281/zenodo.7730978> (Hu et al., 2023).

Supplement. The supplement related to this article is available online at: <https://doi.org/10.5194/acp-23-5517-2023-supplement>.

Author contributions. XH performed data analysis, prepared the figures, and wrote the manuscript. JS designed the experiment and outlined the manuscript. XH, CX, and JS conducted the measurements. XS, YZ, QL, ZL, SZ, JW, AY, JL, SL, and XZ discussed the results and commented on the manuscript.

Competing interests. The contact author has declared that none of the authors has any competing interests.

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